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HAES REPORT NO. 40

ANALYSIS OF HIGH ALTITUDE EFFECTS SIMULATION (HAES)

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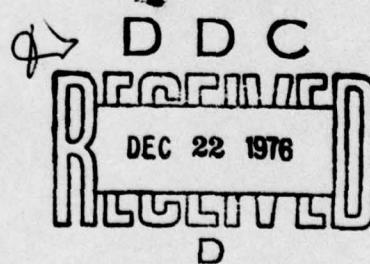
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AIR FORCE GEOPHYSICS LABORATORY
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PREFACE

The High Altitude Effects Simulation (HAES) Program sponsored by the Defense Nuclear Agency since the early 1970 time period, compromises several groupings of separate, but interrelated technical activities, e.g., ICECAP (Infrared Chemistry Experiments- Coordinated Auroral Program). Each of the latter have the common objectives of providing information ascertained as essential for the development and validation of predictive computer codes designed for use with high priority DoD radar, communications, and optical defensive systems.

Since the inception of the HAES Program, significant achievements and results have been described in reports published by DNA, participating service laboratories, and supportive organizations. In order to provide greater visibility for such information and enhance its timely applications, significant reports published since early calendar 1974 shall be identified with an assigned HAES serial number and the appropriate activity acronym (e.g., ICECAP) as part of the report title. A complete and current bibliography of all HAES reports issued prior to and subsequent to HAES Report No. 1, dated 5 February 1974 entitled, "Rocket Launch of an SWIR Spectrometer into an Aurora (ICECAP 72)", AFCRL Environmental Research Paper No. 466, is maintained and available on request from DASIAC, DoD Nuclear Information and Analysis Center, 816 State Street, Santa Barbara, California, 93102, Telephone [805] 965-0551.

This report, which is the first scientific report under AFCRL Contract No. F19628-74-C-0177 is the fortieth report in the HAES series, and covers technical activities period 16 April 1974 through 15 April 1975.

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The calculations in this report provided a portion of the theoretical data base on which the EXCEDE series of rocketborne electron experiments were designed. The support of Dr. Harold C. Fitz, Jr and LCDR Chris Thomas of DNA and Mr. Herb Mitchell of R&D Associates and the technical direction of Dr. A.T. Stair,Jr., Mr. James Ulwick and Mr. Robert R. O'Neil of AFGL is gratefully acknowledged.

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1.0 INTRODUCTION

Contract F19628-74-C-0177 is a contract to analyze High Altitude Effects Simulation (HAES) Experiments performed as part of the ICECAP and EXCEDE Programs and the application of these measurements to physics and chemistry code modelling of atmospheric emissions, such as OPTIR. The work performed under the initial phase of this program includes modelling of energy deposition, particularly for EXCEDE type experiments where radial distributions are probably more important than in auroral events (Section 2.0); calculations of the zenith spectral radiance at 4.3 μm (Section 3.0); and calculations of vibrational population of the first excited level of the CO_2 ν_3 vibrational mode (Section 4.0).

2.0 ENERGY DEPOSITION

The energy deposition calculations are based on the Berger, et al^[1] Monte Carlo results which are briefly summarized in the functional relations

$$F(h, r) = \rho(h) R(h, r) A(z_m)$$

$$R(h, r) = F(h, r) \int_0^\infty F(h, r') 2r' dr'$$

$$A(z_m) = (E/r_p) f(z_m/r_p)$$

$$z_m = \int_h^\infty \rho(h') dh'$$

where $F(h, r)$ is the energy per unit volume deposited at a distance $h-h_0$ down along the field line and r radially outward from a magnetic field line passing through the injection point of a single electron. Needed to calculate this are $\rho(h)$, the atmospheric density at the point of observation; $R(h, r)$ the radial distribution function; z_m , the mass thickness between the points of injection and observation; and r_p , the practical range of the incident electron (Figures 1, 2, 3).

The Monte Carlo calculations have been fitted by Berger, et al, to the following functional approximations:

$$R(h, r) \simeq \frac{1}{2\pi r_H^2} \frac{b^a}{f(a)} \left(\frac{r}{r_H}\right)^{a-2} e^{-b(r/r_H)}$$

where

r_H - the initial Larmor radius of the incident electron

a, b - dimensionless parameters related to moments of $R(h, r)$

where $\langle \rho \rangle$ is the mean radius and the variance is σ^2 ; $\sigma^2 = \langle \rho^2 \rangle - \langle \rho \rangle^2$

$$a = \frac{\langle \rho \rangle^2}{\sigma^2}; b = \frac{a r_H}{\langle \rho \rangle}$$

Figure 4 is a plot of $1/\sqrt{a}$ as a function of z_m/r_p . Figure 5 is a plot of a/b as a function of energy for different values of z_m/r_p . Figure 6 is a plot of a/b as a function of z_m/r_p for 3 keV electrons.

$100 \cdot 10^{-12}$

10^{-11}

$\rho(h) [g/cm^3]$

10^{-10}

10^{-9}

h
(km)
 $z_m(h)$ — mass thickness from
 h to 300 km.

200

180

160

140

120

100

$20 \cdot 10^{-6}$

$z_m(h) [g/cm^2]$

10^{-4}

10^{-3}

10

$\rho(h)$

$z_m(h)$

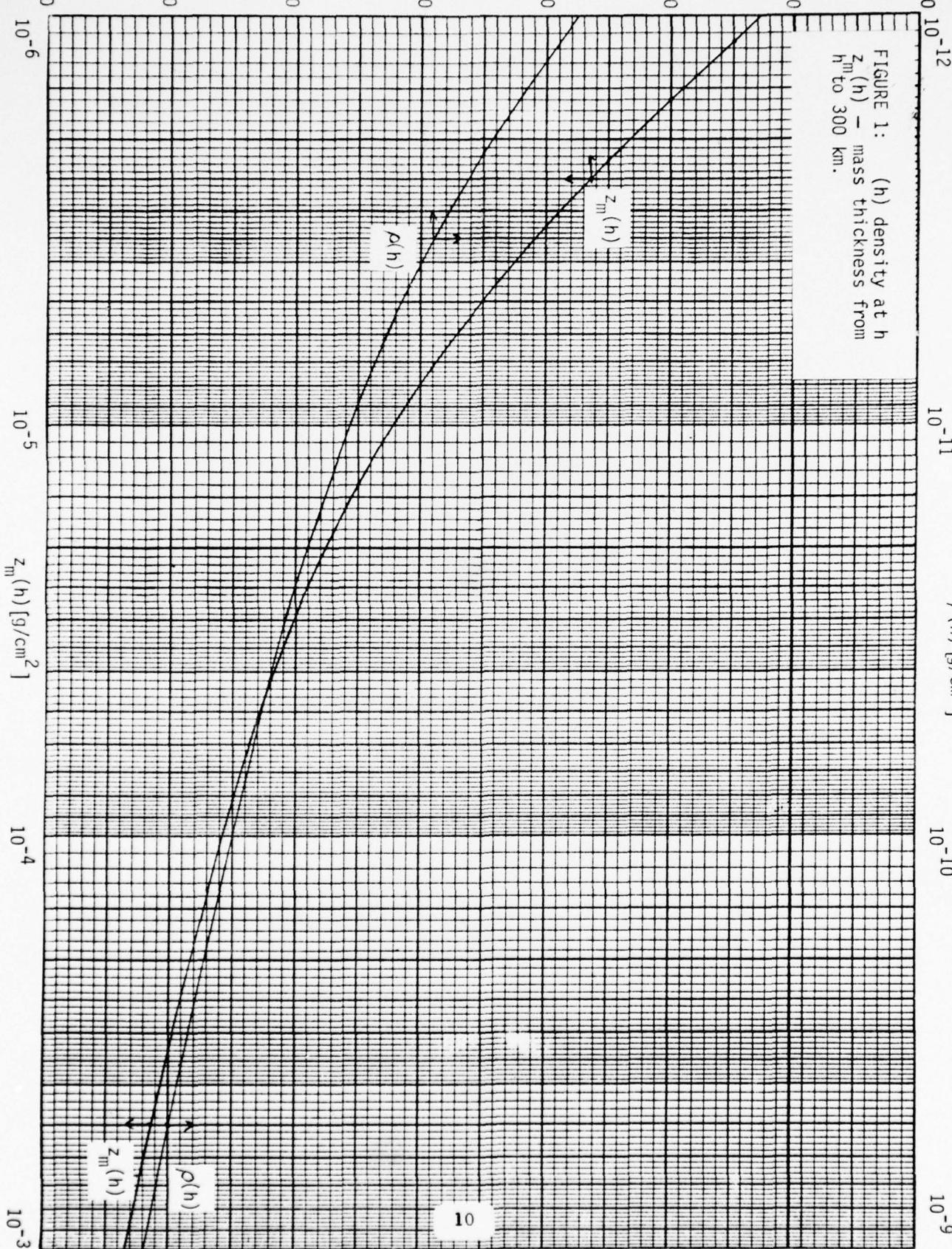


FIGURE 2: Practical Range (r_p) versus Electron Energy (E_0)

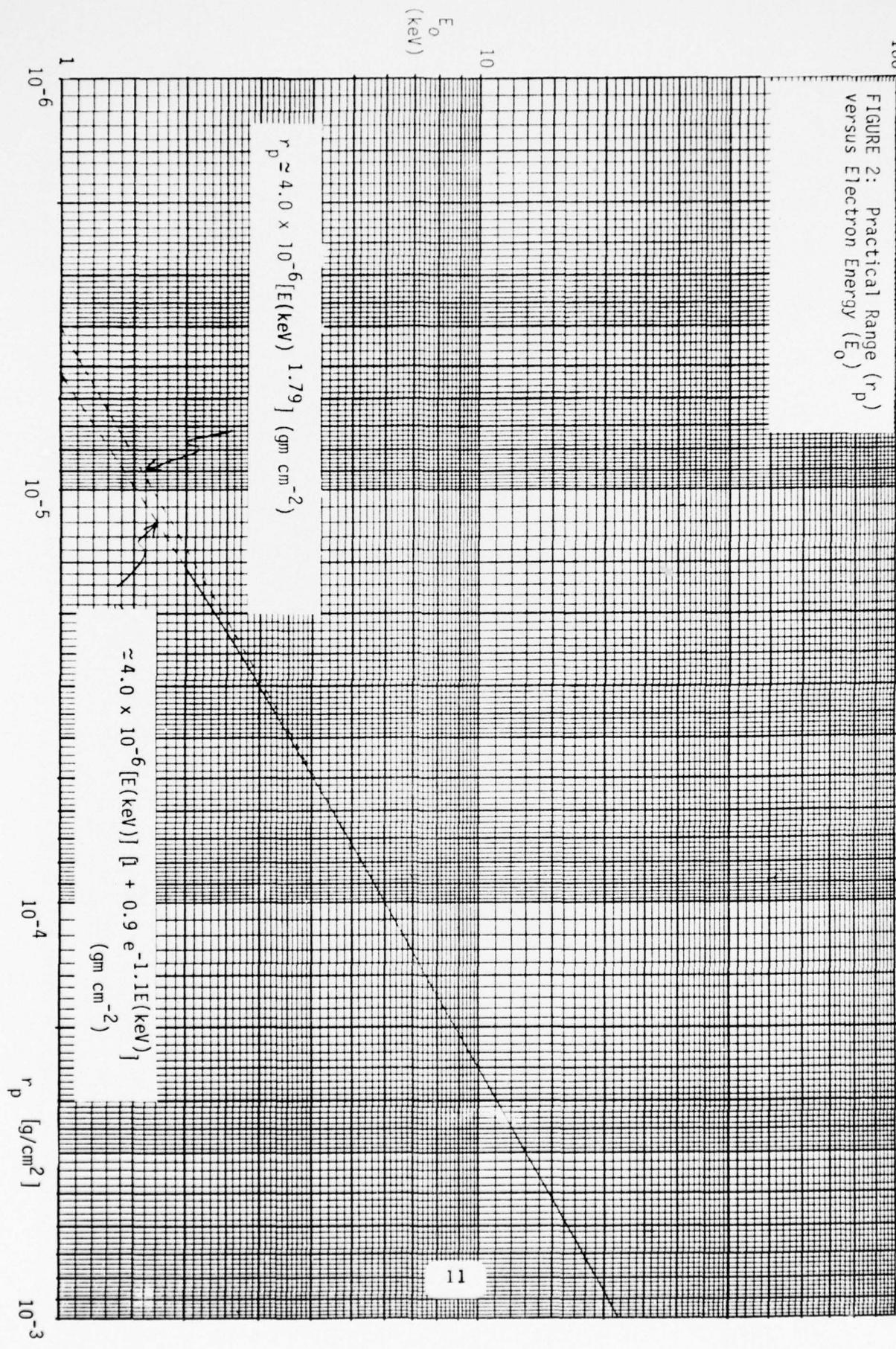
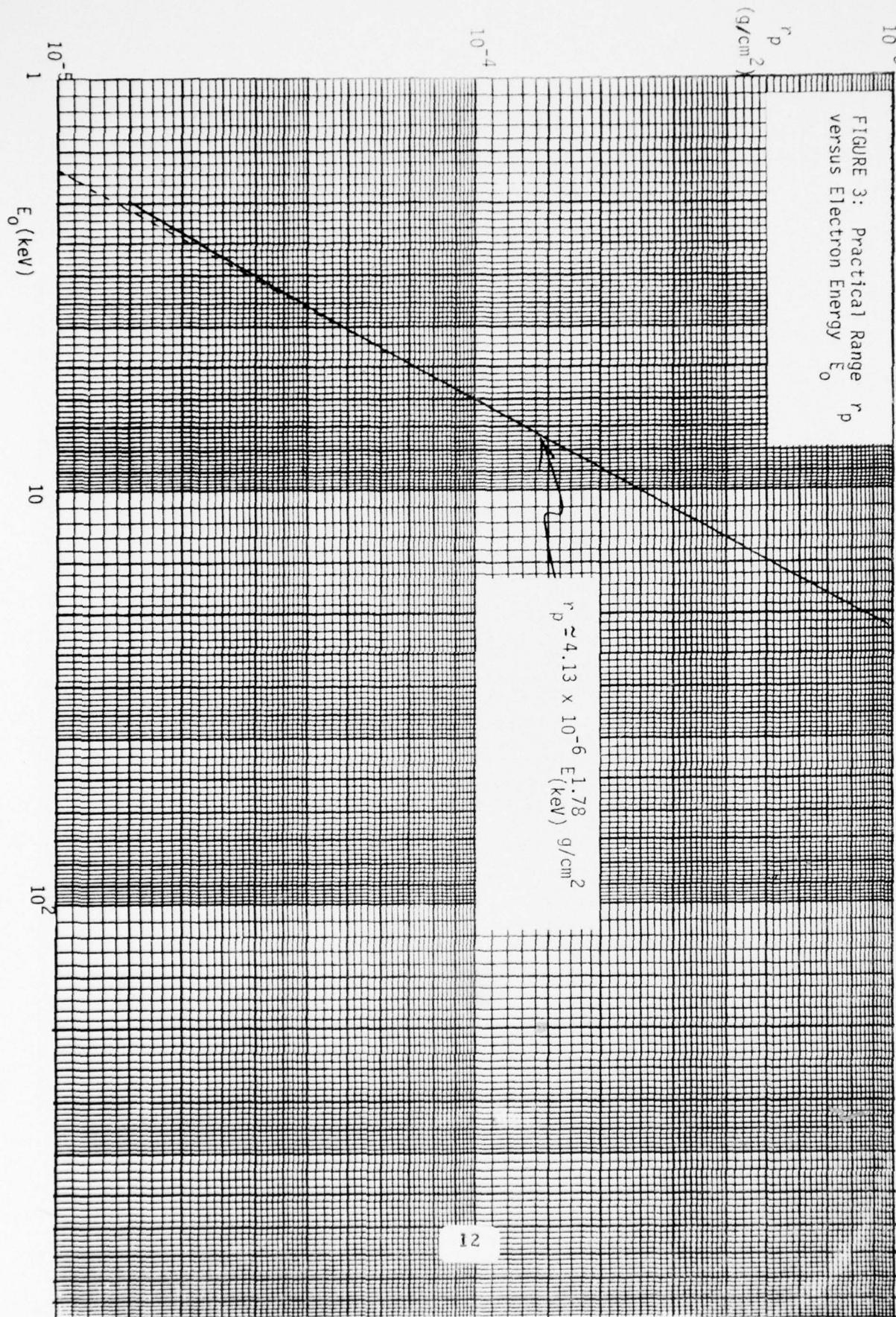
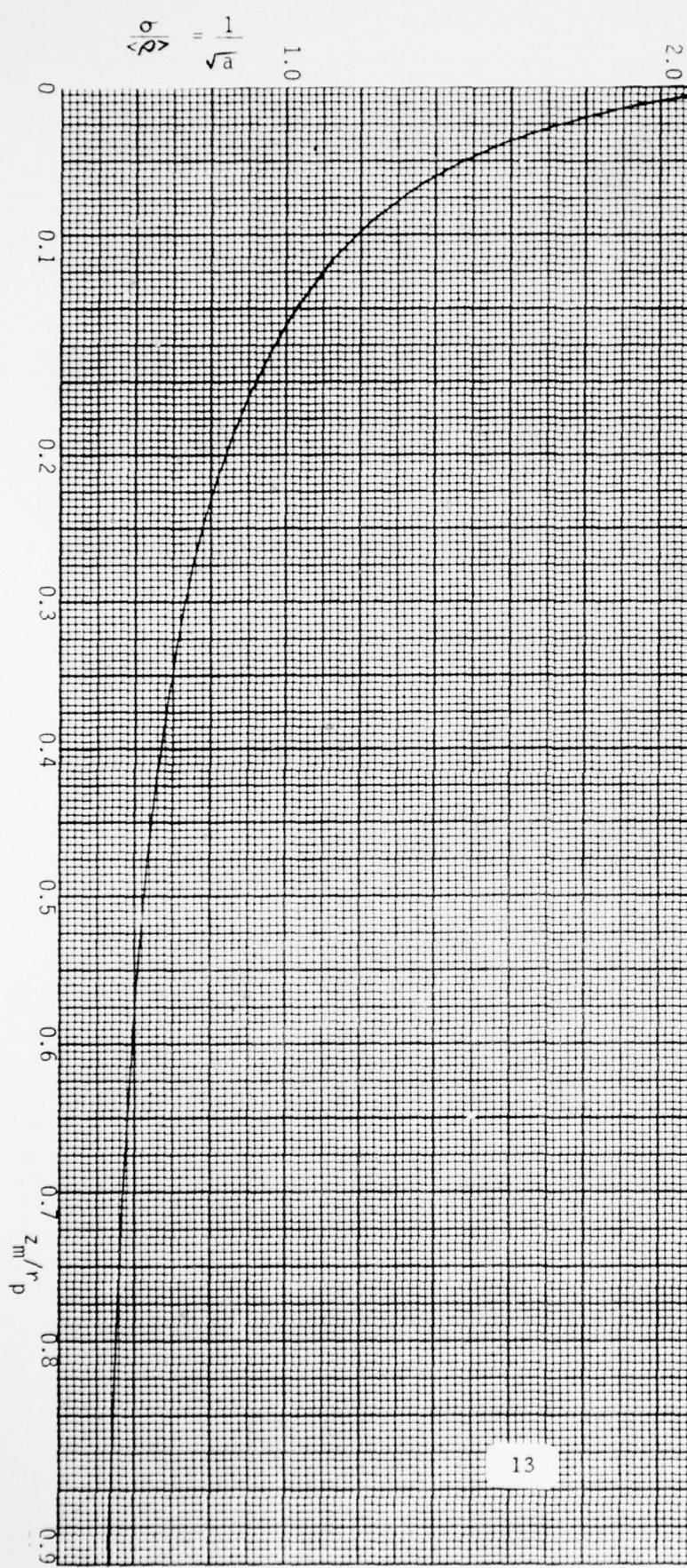
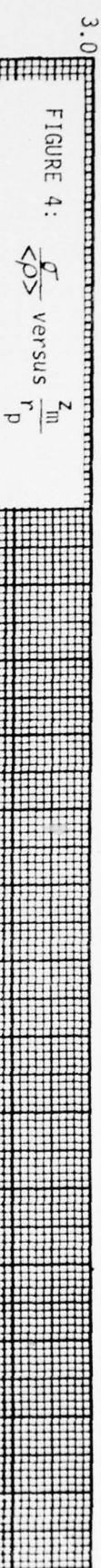


FIGURE 3: Practical Range r_p versus Electron Energy E_0



$2 \leq E_0 \leq 20$ keV $\Theta_0 = 0^\circ$



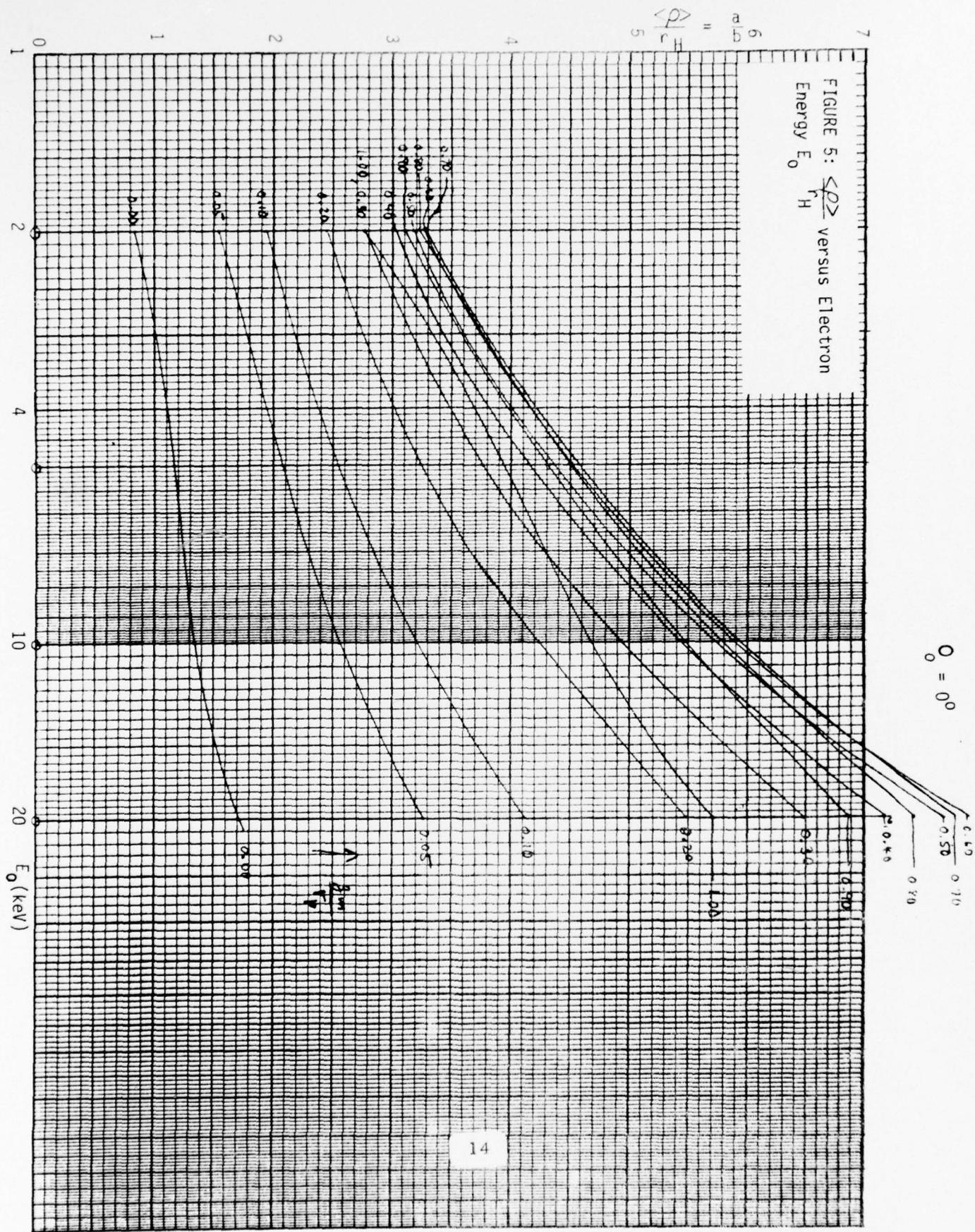
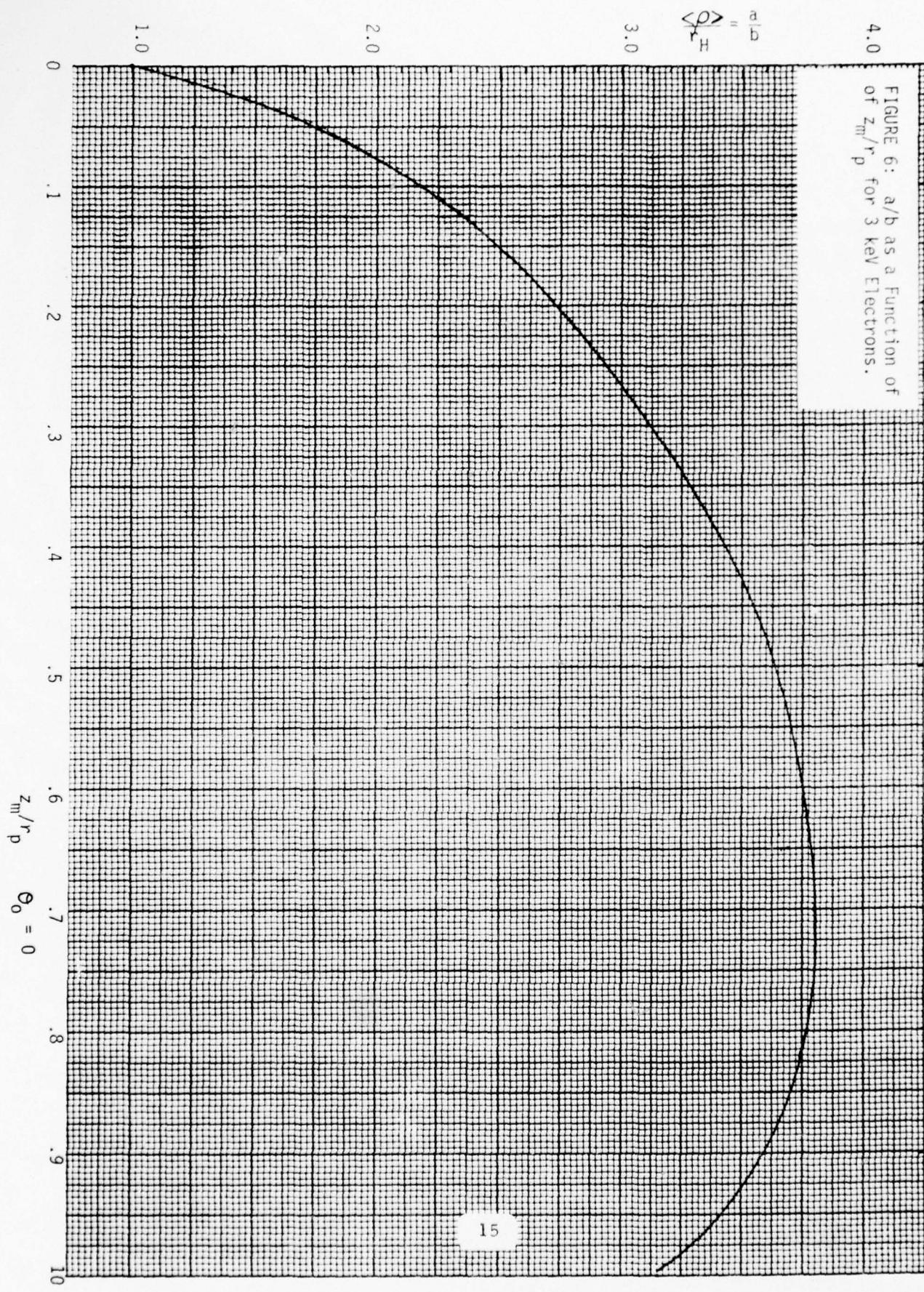


FIGURE 6: a/b as a Function of
of z_m/r_p for 3 keV Electrons.



Tables in the Berger, et al paper provide: radial distribution a and b parameters for a range of energies and distances from the point of injection; atmospheric density as a function of altitude; and values of the "universal" f-function of the axial factor, $A(z_m)$, for various distances from the source. These tabular entries will be least-squares fitted to appropriate functions and incorporated into a subroutine that returns the energy deposition rate and ionization production rate, $Q = \dot{n}_e = F(h, r) \cdot I_o / ew$ (I_o - beam current, e - electron charge, w - eV/ion pair), for a given energy and geometry.

Recombination

For sufficiently high ionization densities, recombination may modify results significantly and thus must be taken into account.

The basic equation relating the ionization density to source strength is:

$$\dot{n}_e(t) = Q(t) - \alpha n_e^2(t)$$

a nonlinear, first-order differential equation of the Riccati type. The analytical solution for arbitrary $Q(t)$ is not known, but through the nonlinear transformation

$$\hat{n}(t) = \exp(\alpha \int_0^t dt' n_e(t'))$$

The linear, but second-order, homogeneous differential equation

$$\hat{\dot{n}}_e(t) - \alpha Q(t) \hat{n}_e(t) = 0$$

is obtained, which for certain source term functional dependences, can be solved in terms of known functions.

Line Integrals

The electron density, $n_e(\vec{r}, t)$, being defined by a nonlinear differential equation involving a time-varying source function, will virtually

never be obtained in an analytical form so that a numerical evaluation of the line integral $\int n_e(\vec{r}, t) dl$, is unavoidable. For each sub-interval of the integral, the value of $n_e(\vec{r}, t)$ will have to be obtained by an integration of the equation $\dot{n}_e(t) = Q(t) - \alpha n_e^2(t)$, where the time variation, $Q(t)$, of the source is due to source trajectory motion and possibly to source intensity changes.

Approximate Calculations

For order of magnitude and qualitative investigations, certain simplified cases have been studied. If recombination during the relatively brief duration of the source pulse, $Q(t)$, can be neglected, then

$$\dot{n}_e(t) = Q(t) - \alpha n_e^2(t) \simeq Q(t)$$

and the electron density during the source pulse is given by

$$n_e(t) \simeq \int_{-\infty}^t dt' Q(t')$$

After the source pulse is over, $\dot{n}_e = -\alpha n_e^2$, which has the solution

$$n_e(t) = \frac{n_0}{1 + \alpha n_0 t}$$

where $n_0 = \int_{-\infty}^{\infty} dt' Q(t')$ is the total electron density produced by the source pulse (in the absence of recombination). Since, for the general point in space,

$$Q(t) = Q(h(t), r(t)) = I_0 F(h, r)/ew$$

where h and r are determined by the (moving) source and observation point geometry, then

$$n_e(\vec{r}, t) = \frac{n_0(\vec{r})}{1 + \alpha n_0(\vec{r})t}$$

where

$$n_e(\vec{r}) = \frac{I_o}{ew} \int_{-\infty}^{\infty} dt' F(h(t), r(t))$$

To perform an analytical line integration of this $n_e(\vec{r}, t)$, the "initial" distribution, $n_o(\vec{r})$, must be of relatively simple algebraic form. Unfortunately, the $n_o(\vec{r})$ from the Berger theory is generally a non-elementary integral, effectively making a closed solution impossible. However, in the special case of line integrals along directions normal to the plane of the magnetic field and source velocity, and at altitudes for which the radial distribution, $R(h, r)$, has an integer a -parameter, $n_o(\vec{r})$ can be evaluated in terms of the zero and first-order modified Hankel functions. These special $n_o(\vec{r})$ cases can be evaluated from tables and least-squares fit to functions that permit analytical evaluation of the line integral to yield the approximate time dependence of $\int n_e dl$ at the special altitudes.

The model distributions which were fit to the special (evaluable) integrals were rectangular (uniform), triangular, exponential and Lorentzian. The resulting time profiles differed by 40% at one characteristic time, τ_o ; by 3:1 at 10 τ_o and by 10:1 at 100 τ_o .

More Exact Calculations

With some idea of the source function pulses, $Q(t)$, to be encountered, the electron density $n_e(\vec{r}, t)$ can be estimated, now taking recombination into account, either by numerically integrating the basic equation

$$\dot{n}_e(t) = Q(t) - \alpha n_e^2(t)$$

defining $Q(t)$ directly in terms of the Berger function, $F(h, r)$, or by segmental representation of $Q(t)$ with polynomial fits to selected values of the Berger function.

There exist a number of straightforward, first-order differential equation numerical integration techniques, but good numerical accuracy

requires evaluation of a relatively complicated $Q(t)$ function many times over. Since the many such time-integrations required in the various desired line integrals might prove too costly in computer time, a linear segment representation approach was also looked into. If the $Q(t)$ function can be adequately represented by five or ten linear segments, then the solution throughout any segment can be expressed in terms of the solutions of the transformed, second-order differential equation:

$$\hat{n}_e(t) + \alpha Q(t) \hat{n}_e(t) = 0$$

where

$$Q(t) = Q_0(1 + t/\tau)$$

the solutions of which are directly related to the Airy integrals $Ai(x)$ and $Bi(x)$ and their derivatives. Subroutines to evaluate these functions for $-\infty < x < \infty$ have been written and await testing.

The energy deposition calculations discussed above are applicable to both the ICECAP and EXCEDE Programs and will be used to evaluate electron spectra such as that presented in Appendix 1. The accuracy of the radial description is more relevant to the EXCEDE geometry than to a typical auroral geometry.

3.0 SPECTRAL RADIANCE CALCULATIONS AT 4.3 μM , 9.6 μM , 15 μM .

The night time zenith spectral radiance has been calculated for the CO₂ bands at 4.3 μm and 15 μm and for the ozone band at 9.6 μm at altitudes for 60 to 120 μm . The spectral radiance was calculated assuming the U.S. Standard Atmosphere (180° K Mesospheric temperature) and an atmosphere with a 230° K Mesospheric temperature. The results are presented in Appendix 2. Degges 1974 Infrared Radiance Model² was used for these calculations.

Degges' work is an extension of the study of Corbin, et al. (1969)³ and Degges (1972).⁴ The former investigated the natural infrared background of the earth in the 5 to 25 micrometer spectral region. For convenience, their study divided the atmosphere into two regions with a division at 70 km. Below 70 km the atmosphere was assumed to be in thermal equilibrium. Above 70 km explicit calculations were made of processes which excite and de-excite molecular vibrational and rotational levels which are the source of infrared radiation. Their study concentrated on radiation from water vapor, carbon dioxide, ozone, nitric oxide and nitrous oxide, which are the principal radiating species in the spectral region considered. In addition, nitric acid was included in the lower atmosphere work and estimates were made of radiation to be expected from particulate matter suspended in the atmosphere.

Corbin, et al. (1969) presented models for the lower atmosphere for a wide range of seasonal and latitudinal conditions. This was not possible for the abundances of most minor neutral species. More data has since become available, particularly for nitric oxide and the hydroxyl radical, but at present it appears that the best means of estimating abundances of important infrared emitting species is chemical rate equation integrations including molecular diffusion and eddy mixing. Even calculations involving transport properties cannot always be accepted because the values of eddy mixing coefficients are to a large extent only informed

guesses and published calculations often use outdated rate coefficients.

Degges (1972) reported a computer program with which to determine diurnal variations in abundances of minor species, to investigate the effects of changes in assumed eddy mixing coefficients, and to estimate the effects of new determinations of chemical rate coefficients.

A second area of study involved in improving the radiance model includes the physical processes that control the population of infrared emitting states of atmospheric molecules. Except for the pure rotational radiation from molecules such as water, the degree of excitation of vibrational levels determines the radiation from infrared emitting molecules. The most important mechanisms are collisional excitation and de-excitation and absorption and re-emission of electromagnetic radiation.

In the troposphere and lower stratosphere, collisional processes are rapid enough to control the population of vibrational levels. Above 30 to 50 km, however, collisional excitation becomes less efficient and radiative processes become important. The combined effects of collisional and radiative processes must therefore be considered. Below an altitude of about 90 km molecular nitrogen and oxygen are the most important collision partners. Above that altitude, atomic oxygen becomes important, both in exciting nitric oxide and in determining molecular oxygen and nitrogen vibrational temperatures.

In determining the effects of radiation on the populations of vibrational levels of infrared active molecules, it is necessary to separate the radiation of a single change in vibrational quantum numbers from the rest of the radiation field. The previously reported studies did this by assuming a Doppler line shape for the individual rotational lines of a band and were able to obtain adequate numerical approximations for radiative transfer functions appropriate to single bands of linear molecules, and less accurately, for water vapor and ozone bands lying in the spectral region of interest.

4.0 MODEL CALCULATIONS OF THE VIBRATIONAL POPULATION OF THE FIRST EXCITED LEVEL. ν_3 VIBRATIONAL MODEL.

In order to gain insight as to possible effects of the time and space variations in auroral excitation of the carbon dioxide (CO_2) infrared band systems, we have continued model calculations of the vibrational population of the first excited level of the $\text{CO}_2 \nu_3$ vibrational mode. The model atmosphere used contains only CO_2 , molecular nitrogen, and atomic and molecular oxygen. The CO_2 mixing ratio is assumed constant at 320 parts per million, by volume. Calculations are restricted to levels two kilometers apart lying between 50 and 150 kilometers altitude. The temperature falls from 270 K at 50 km to a mesopause temperature of 191 between 80 and 90 km. Above 90 km, the temperature rises to 661 K at 150 km.

The radiative transfer model of Degges (1974) is used with the collision processes of Table 1, the rates shown being those given by Garvin and Hampson (1974).⁵ The QCHEM integration algorithm described by Manley, et al (1973)⁶ is used to integrate rate equations for the vibrational populations of the $\text{CO}_2 \nu_3$ mode and the first vibrational levels of N_2 and O_2 . The results presented in the following figures should be regarded as only semi-quantitative. The system of rate equations is stiff with coupling between stiff components through the radiation field. Manley, et al (1973) point out that in cases of this nature, the solution given by the QCHEM algorithm relaxes in the correct direction, but the numerical results are not properly correlated with the prescribed time steps.

Results are presented here for two (2) sample cases. In the first, vibrational temperatures at all altitudes were set to 250 K initially. Assuming night time conditions, the system was integrated for 10^6 seconds. Then, an auroral "drizzle" was introduced with a peak ionization rate of 2.5×10^3 ion pairs $\text{cm}^{-3} \text{ sec}^{-1}$ at 110 km and the integration continued for 10^6 seconds. Then, a strong arc was turned on, with a peak ionization

TABLE 1

Collisional De-excitation Rates

1. $N_2 (v = 1) + M \rightarrow N_2 (v = 0) + M$
 $k_1 = 8.53 \times 10^{-7} \exp (-273.10/T^{1/3})$
2. $O_2 (v = 1) + M \rightarrow O_2 (v = 0) + M$
 $k_2 = 4.81 \times 10^{-8} \exp (-169.60/T^{1/3})$
3. $CO_2 (001) + N_2 (v = 0) \rightarrow CO_2 (000) + N_2 (v = 1)$
 $k_3 = 1.71 \times 10^6 \exp (-175.30/T^{1/3}) + 6.07 \times 10^{-14} \exp (15.27/T^{1/3})$
4. $N_2 (v = 1) + O_2 (v = 0) \rightarrow N_2 (v = 0) + O_2 (v = 1)$
 $k_4 = 1.74 \times 10^{-10} \exp (-124.00/T^{1/3})$
5. $N_2 (v = 1) + O \rightarrow N_2 (v = 0) + O$
 $k_5 = 1.07 \times 10^{-10} \exp (-69.90/T^{1/3})$
6. $O_2 (v = 1) + O \rightarrow O_2 (v = 0) + O$
 $k_6 = 6.88 \times 10^{-9} \exp (-76.65/T^{1/3})$

rate of 10^6 ion pairs $\text{cm}^{-3} \text{ sec}^{-1}$ at 100 km and the integration continued for 10^6 seconds. Following this, the auroral excitation was turned off and relaxation followed for 10^4 seconds. The procedure was repeated for an initial vibrational temperature of 300 K. In both cases, the assumption was made that for each ion pair produced, three nitrogen molecules are excited to the first vibrational level.

Figures 1 through 4 display results for the case where initial vibrational temperatures are 250 K. Figures 5 through 8 display results for the case where initial vibrational temperatures are 300 K.

In Figure 1, vibrational populations increase with time below 60 km as collisions and radiative transport are effective in bringing vibrational temperatures closer to kinetic temperatures. Above 75 km, populations fall with time. At late times, populations rise above 125 km as collisions are effective in increasing nitrogen vibrational temperatures toward kinetic temperatures, with subsequent transfer of vibrational excitation to CO₂.

There is little change in Figure 2, the "drizzle" excitation not being large compared to the normal collisional and radiative mechanisms.

The auroral effects are clearly evident in Figure 3. Above 100 km, a balance is reached between enhanced auroral excitation and radiative decay by 10^4 seconds. This aurora produces no ionization at 90 km and below. At late times, vibrational populations continue to increase at altitudes as low as 85 km as energy is transported down by radiation. This downward transport of radiation continues after the aurora is turned off, as shown in Figure 4.

The behavior for the higher initial vibrational temperature of 300 K is similar, as shown in Figures 5 through 8, except that vibrational temperatures fall at all times for altitudes below 65 km. The auroral effects are not as visibly pronounced, because of the higher initial vibrational temperature.

Figures 9 through 12 present vibrational temperatures corresponding to the vibrational populations of Figures 1 through 4.

Figure Captions

Figure 1 $\text{CO}_2 \nu_3$ vibrational populations at $10^{-4}, 10^{-3}, \dots 10^5, 10^6$ seconds following relaxation from a uniform vibrational temperature of 250 K.

Figure 2 $\text{CO}_2 \nu_3$ vibrational populations at $10^{-4}, 10^{-3}, \dots 10^5, 10^6$ seconds with auroral "drizzle", following conditions of Figure 1.

Figure 3 $\text{CO}_2 \nu_3$ vibrational populations at $10^{-4}, 10^{-3}, \dots 10^5, 10^6$ seconds with strong auroral arc, following conditions of Figure 2.

Figure 4 $\text{CO}_2 \nu_3$ vibrational populations at $10^{-6}, 10^{-5}, \dots 10^3, 10^4$ seconds during relaxation from conditions of Figure 3.

Figure 5 $\text{CO}_2 \nu_3$ vibrational populations at $10^{-4}, 10^{-3}, \dots 10^5, 10^6$ seconds following relaxation from a uniform vibrational temperature of 300 K.

Figure 6 $\text{CO}_2 \nu_3$ vibrational populations at $10^{-4}, 10^{-3}, \dots 10^5, 10^6$ seconds with auroral "drizzle", following conditions of Figure 5.

Figure 7 $\text{CO}_2 \nu_3$ vibrational populations at $10^{-4}, 10^{-3}, \dots 10^5, 10^6$ seconds with strong auroral arc, following conditions of Figure 6.

Figure 8 $\text{CO}_2 \nu_3$ vibrational populations at $10^{-6}, 10^{-5}, \dots 10^3, 10^4$ seconds during relaxation from conditions of Figure 7.

Figure 9 $\text{CO}_2 \nu_3$ vibrational temperatures corresponding to vibrational populations of Figure 1.

Figure 10 $\text{CO}_2 \nu_3$ vibrational temperatures corresponding to Figure 2.

Figure 11 $\text{CO}_2 \nu_3$ vibrational temperatures corresponding to Figure 3.

Figure 12 $\text{CO}_2 \nu_3$ vibrational temperatures corresponding to Figure 4.

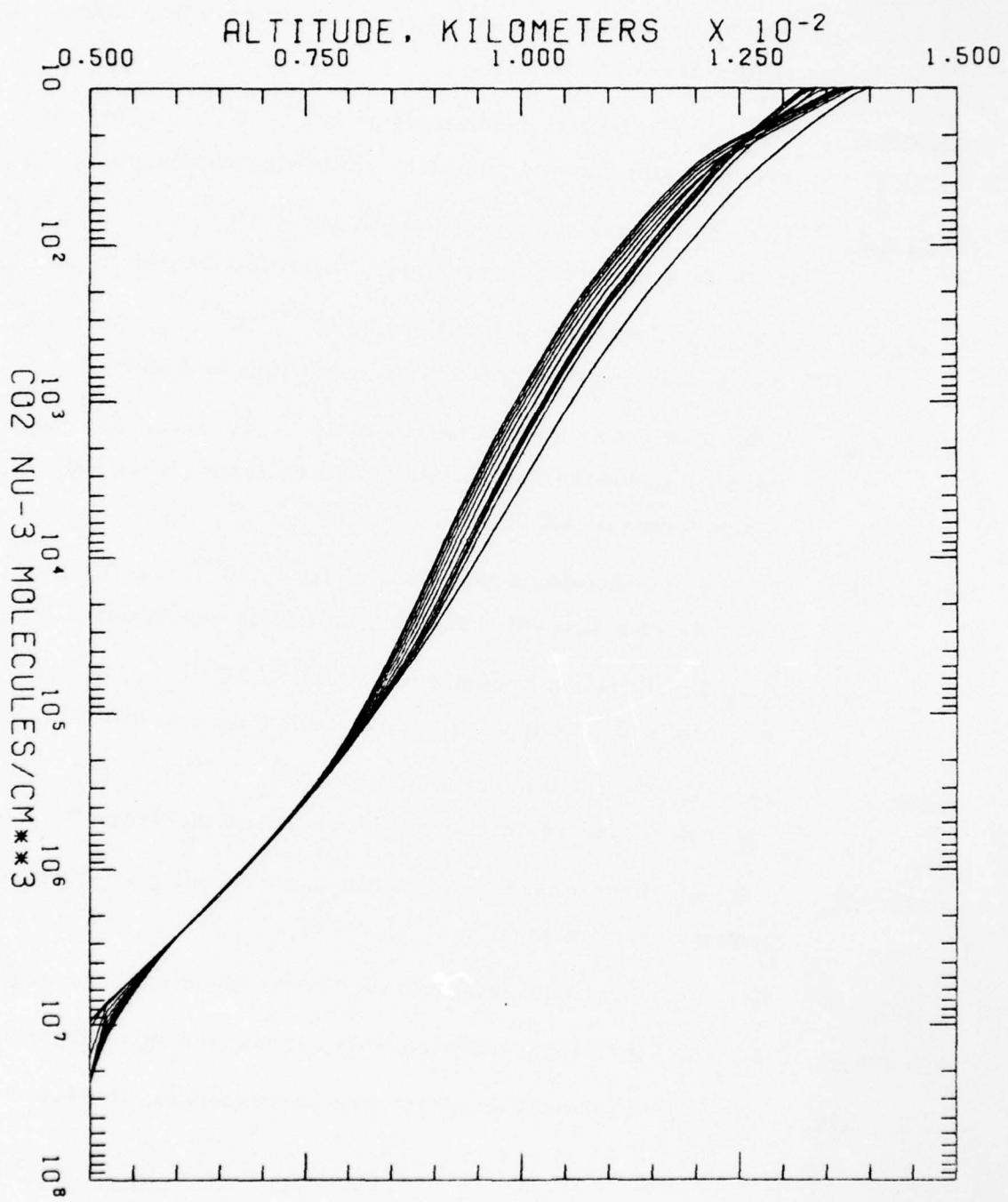


Figure 1

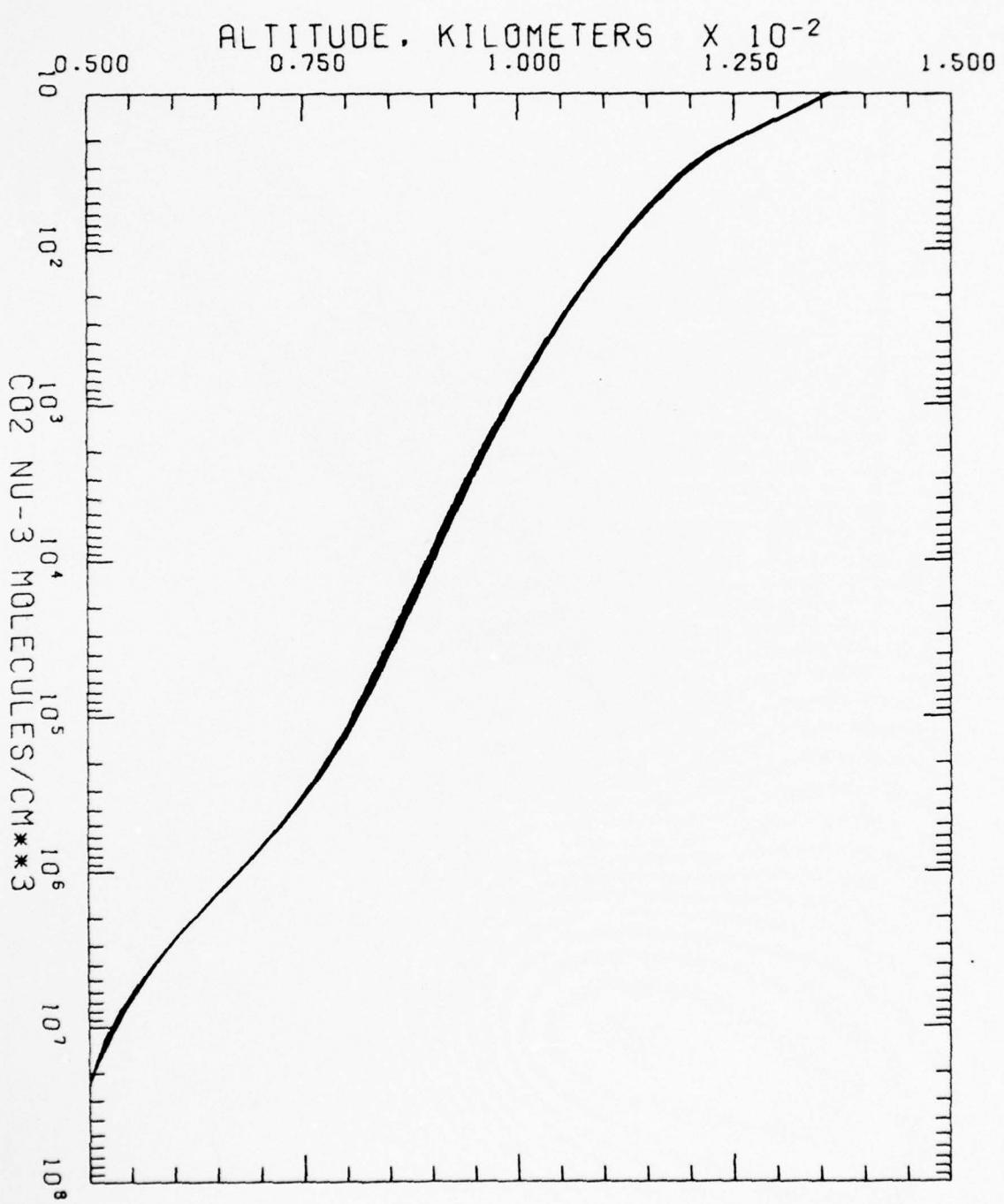


Figure 2

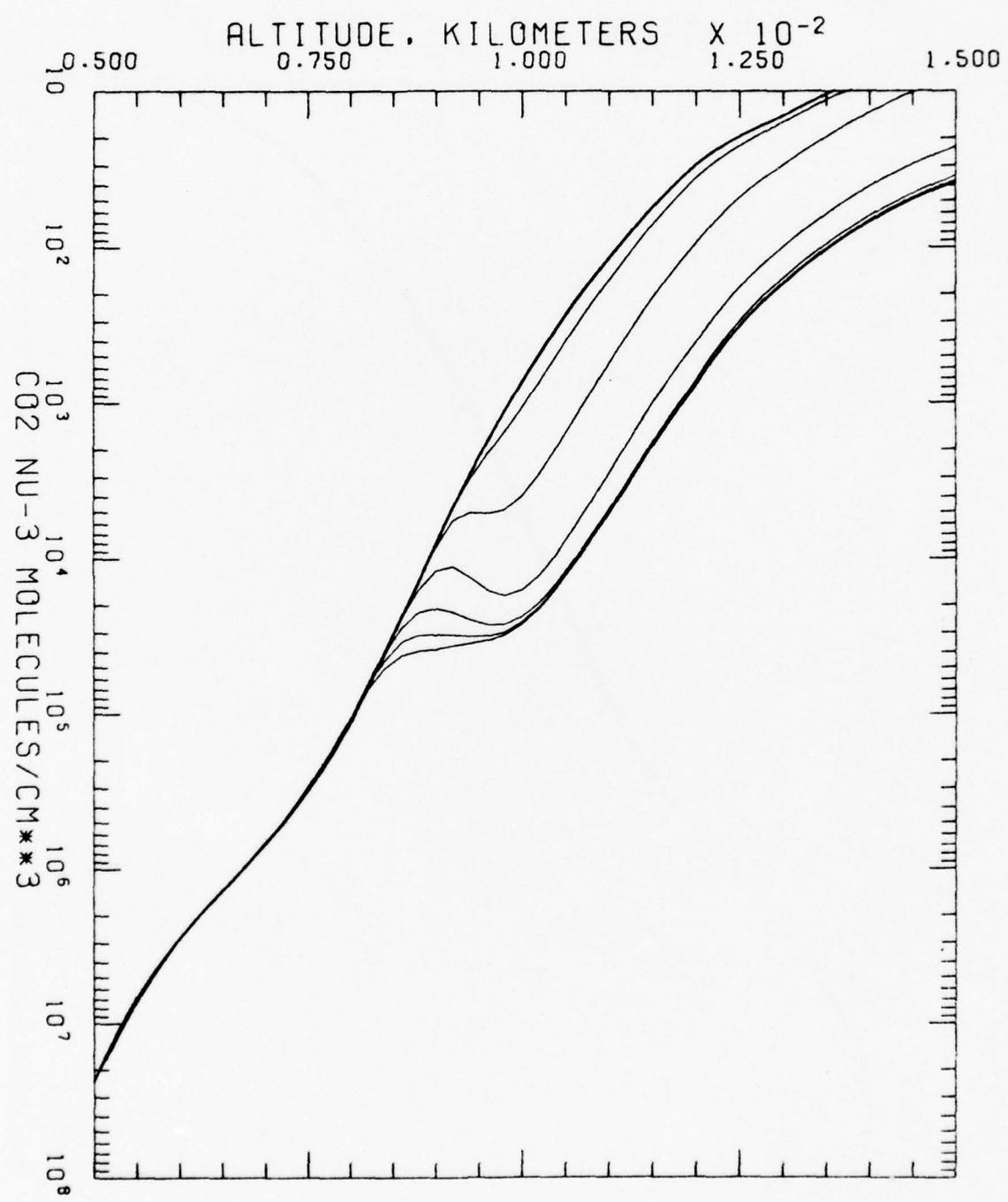
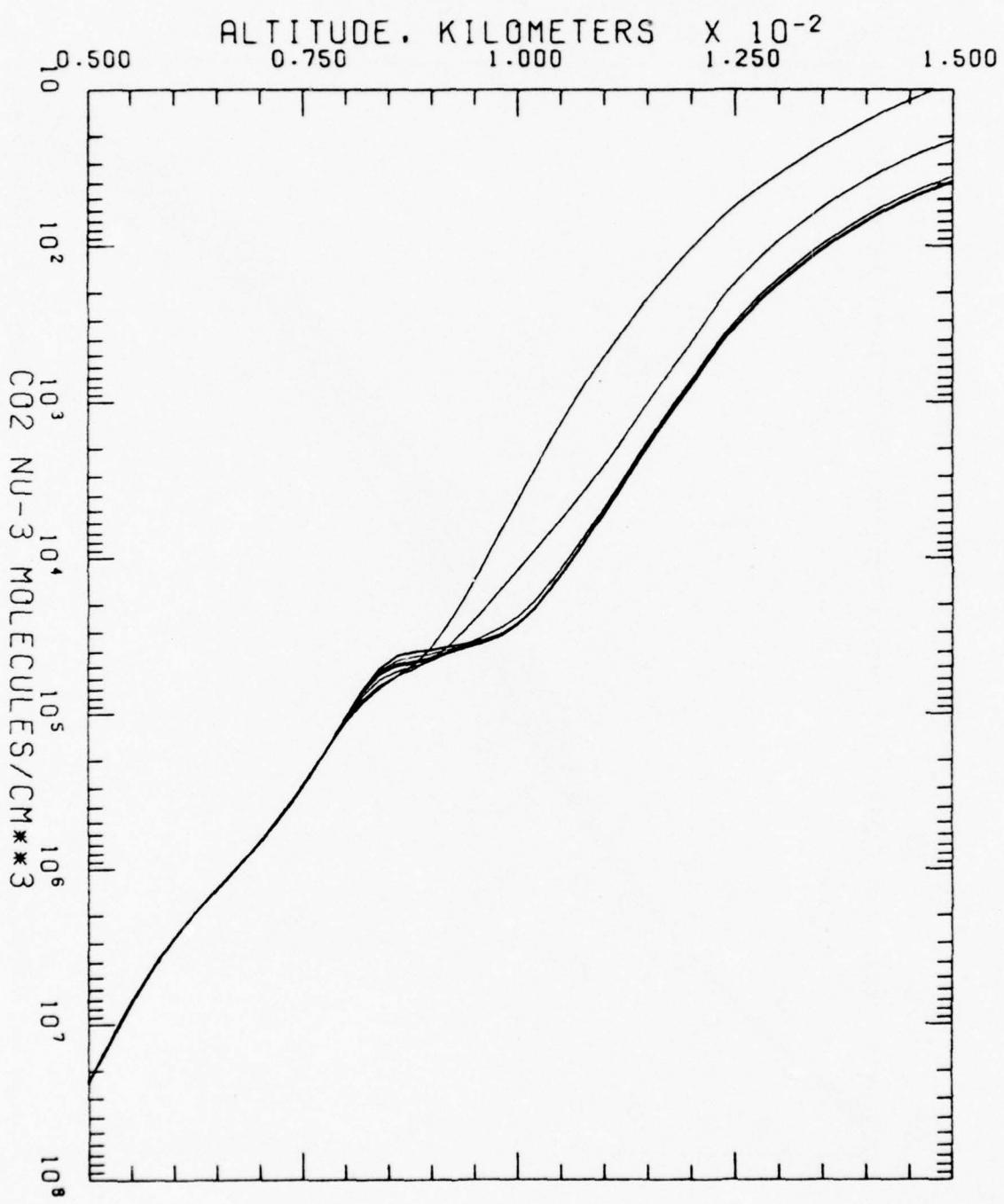


Figure 3



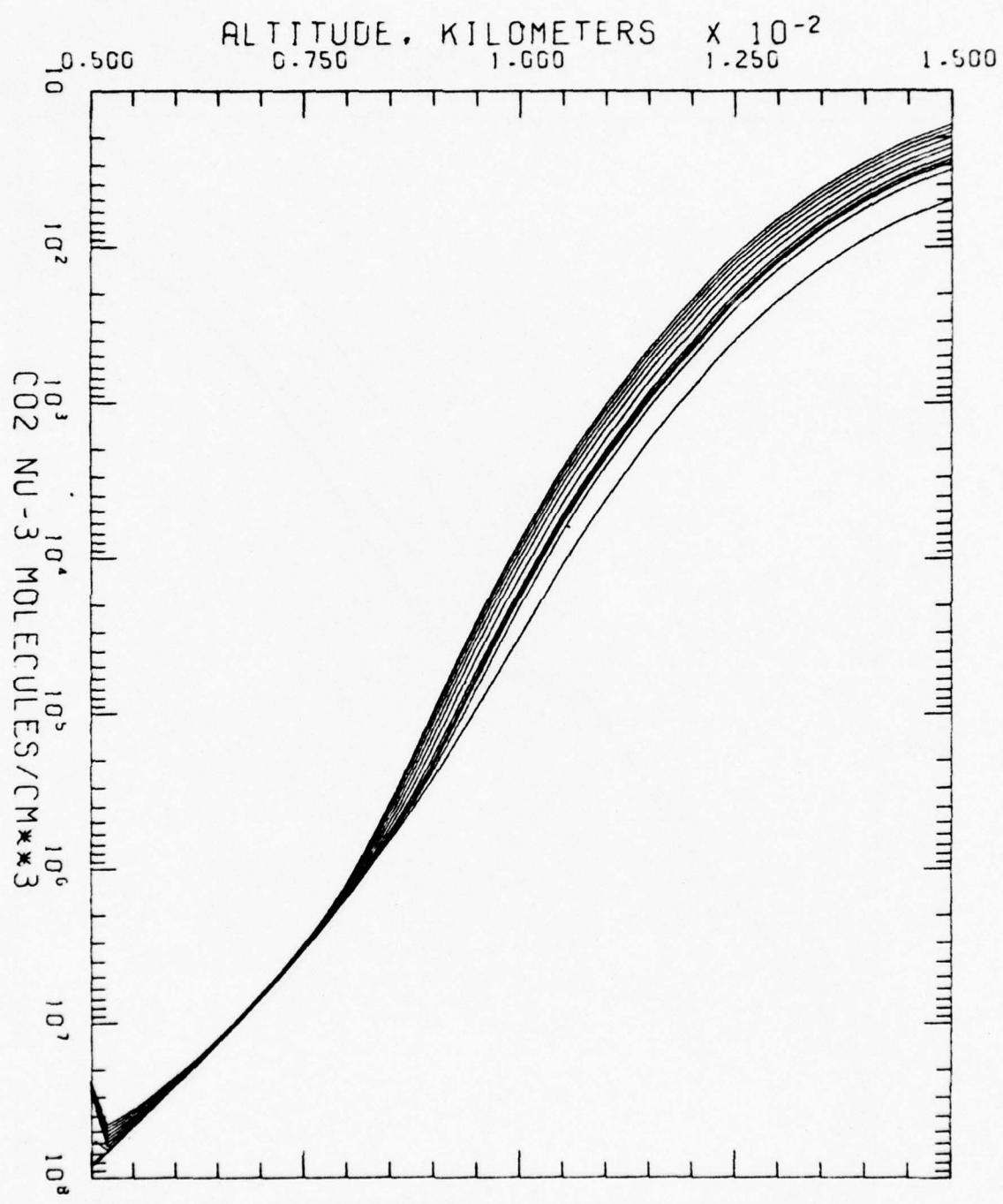


Figure 5

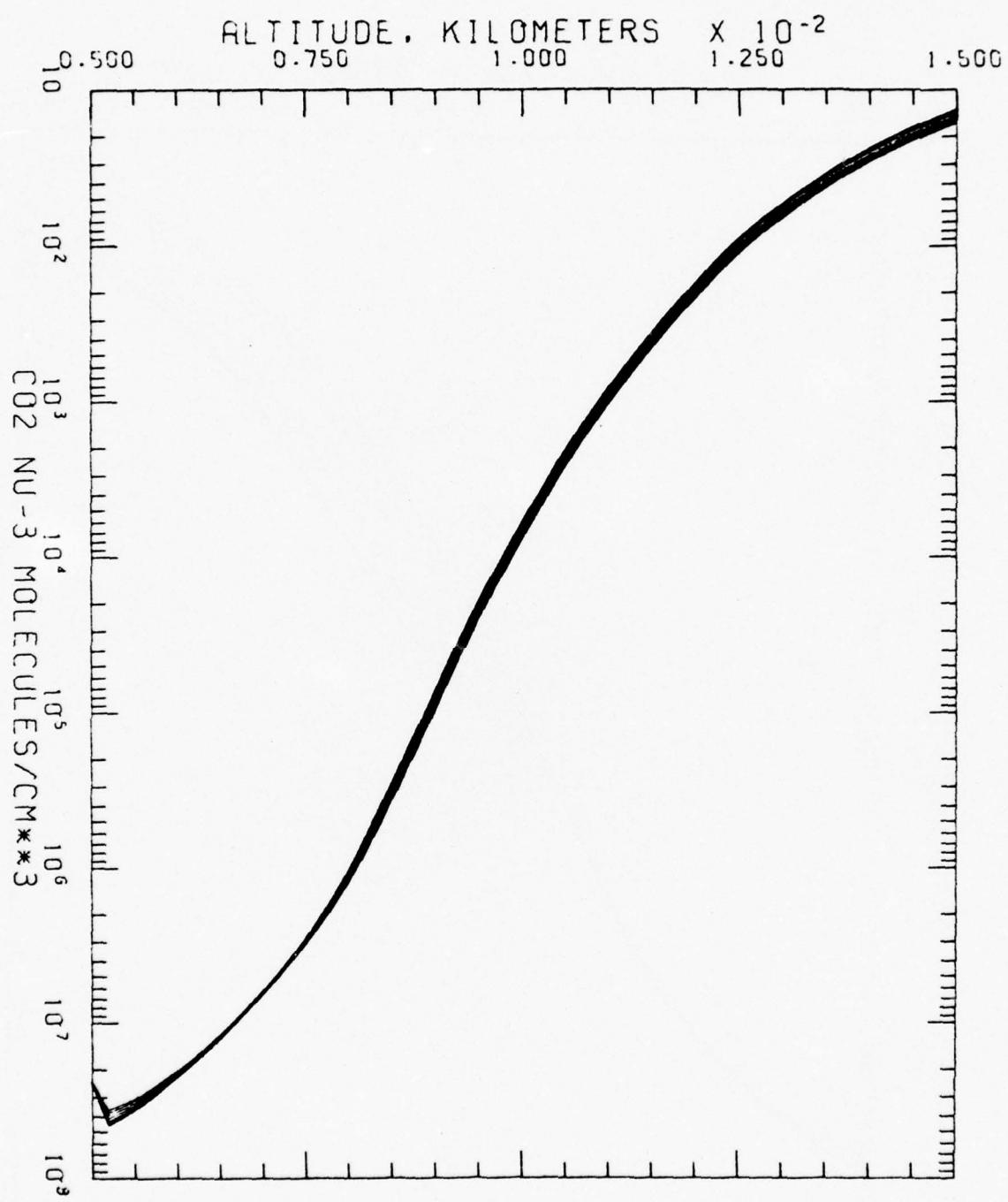


Figure 6

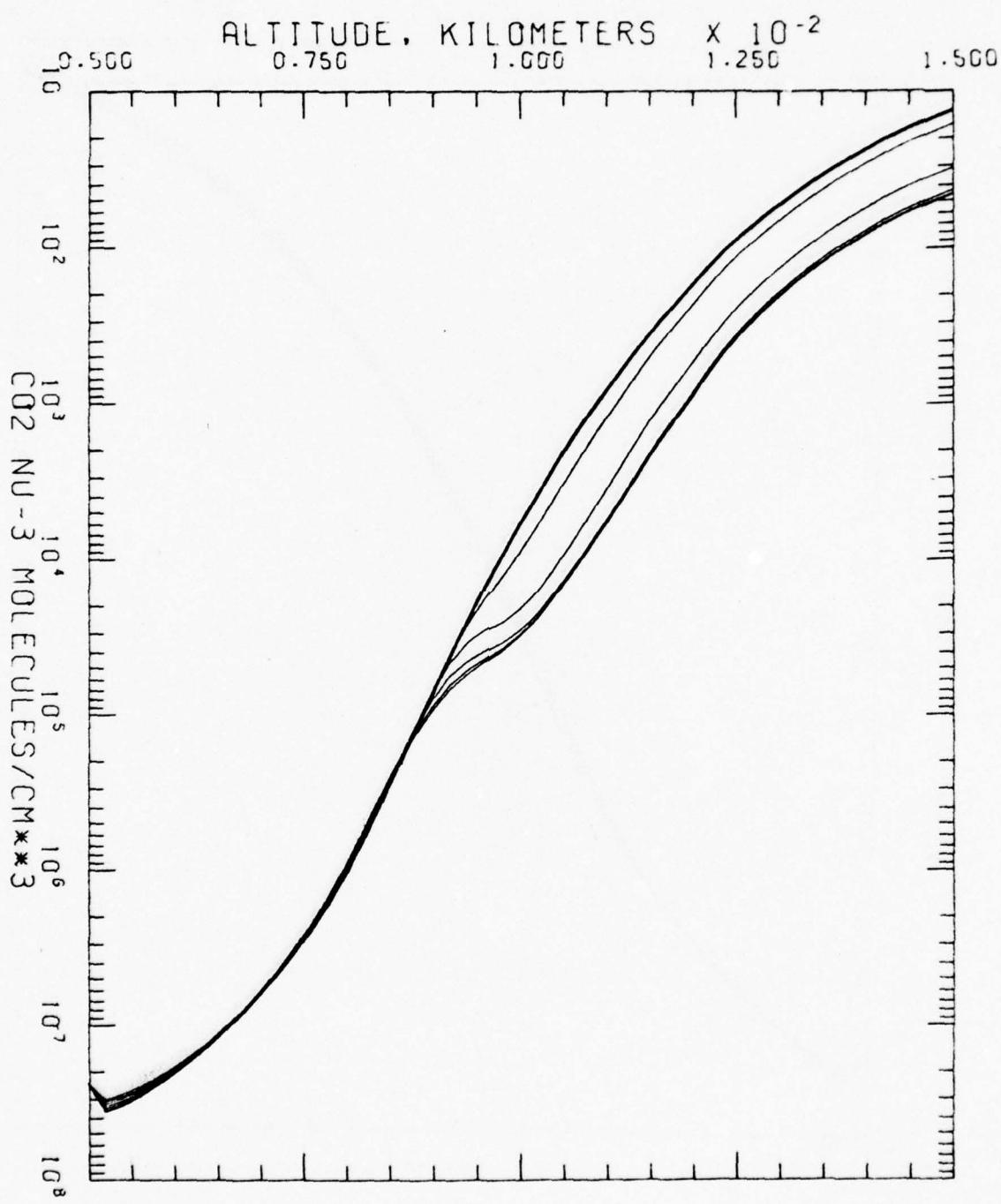


Figure 7

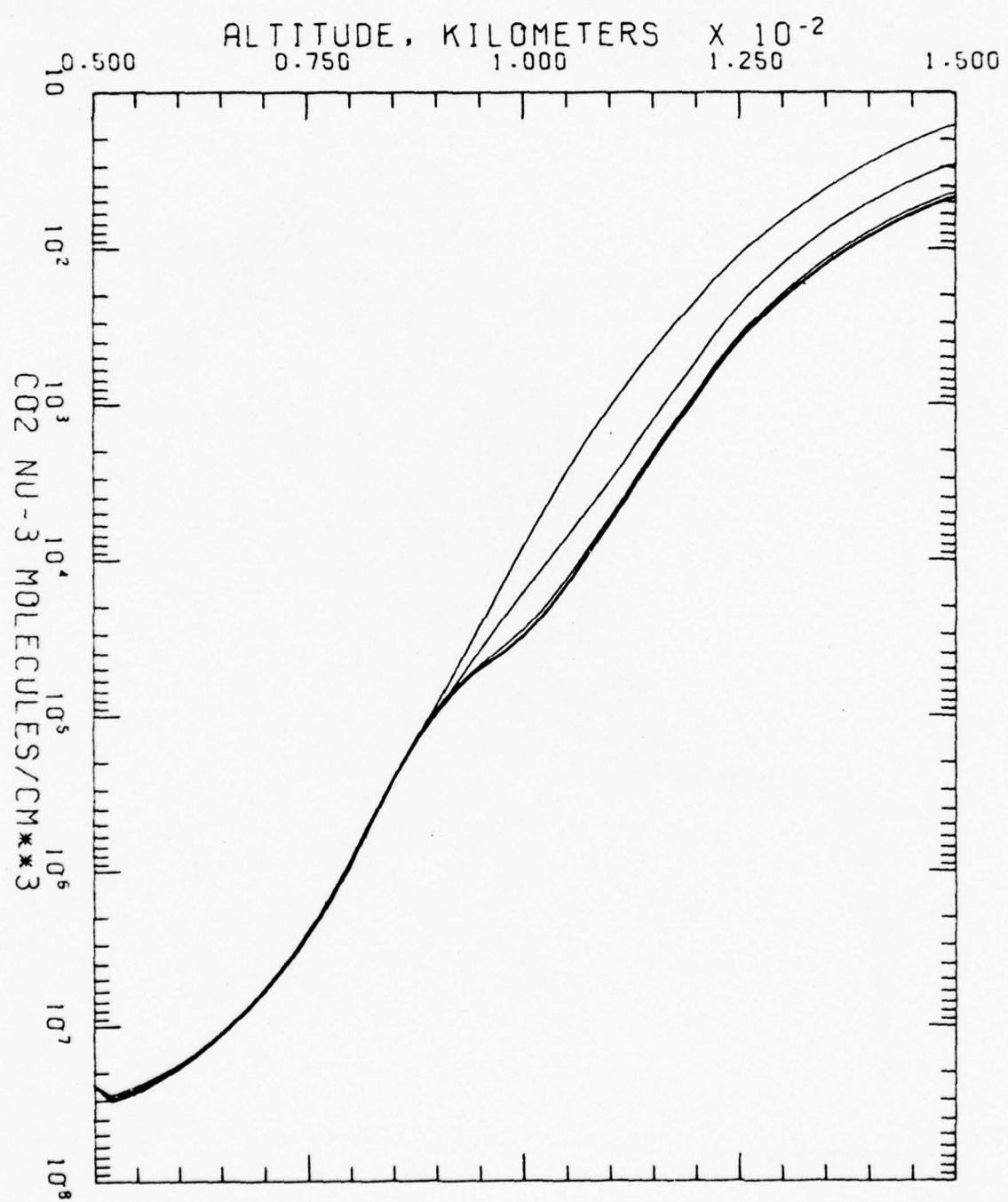


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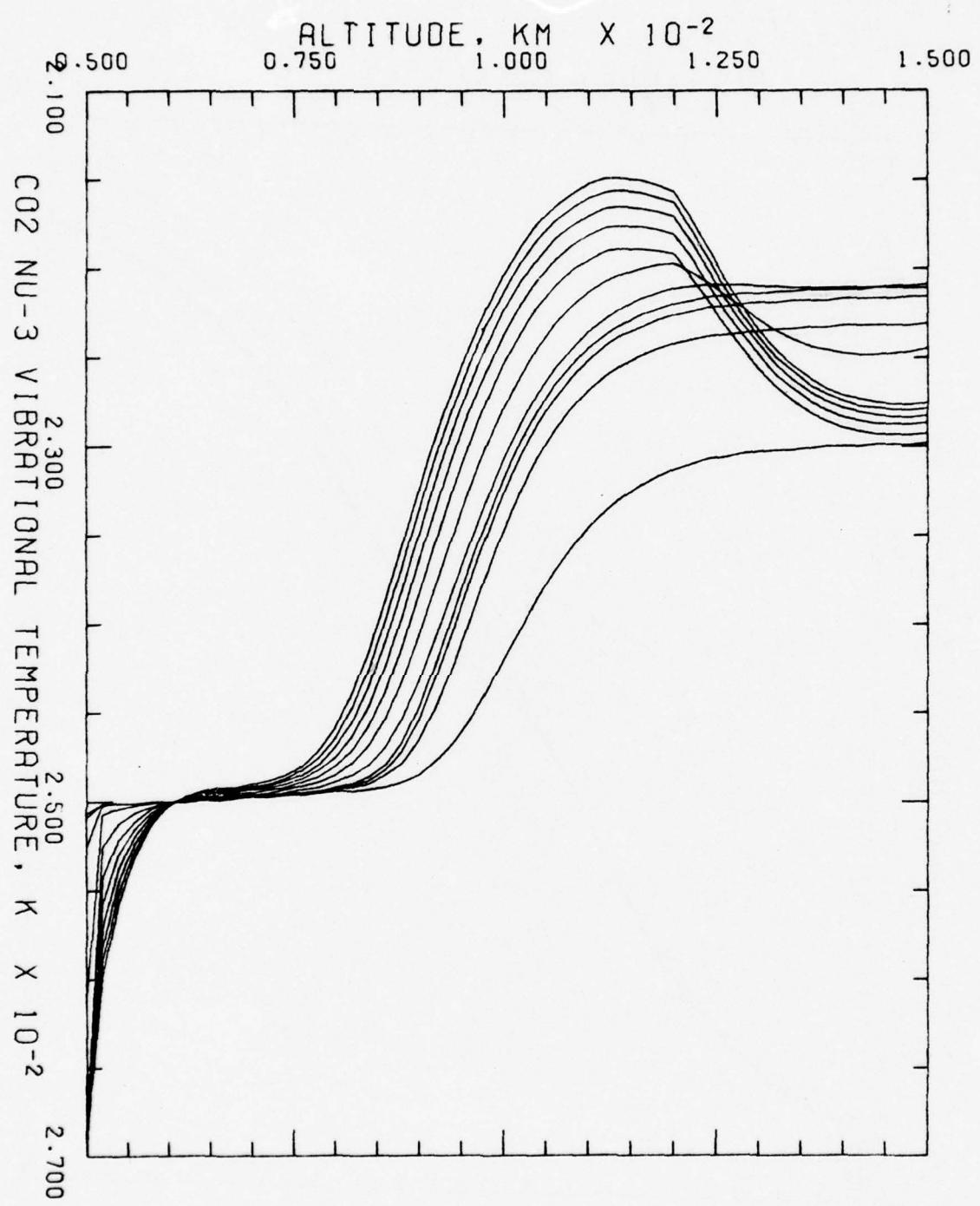


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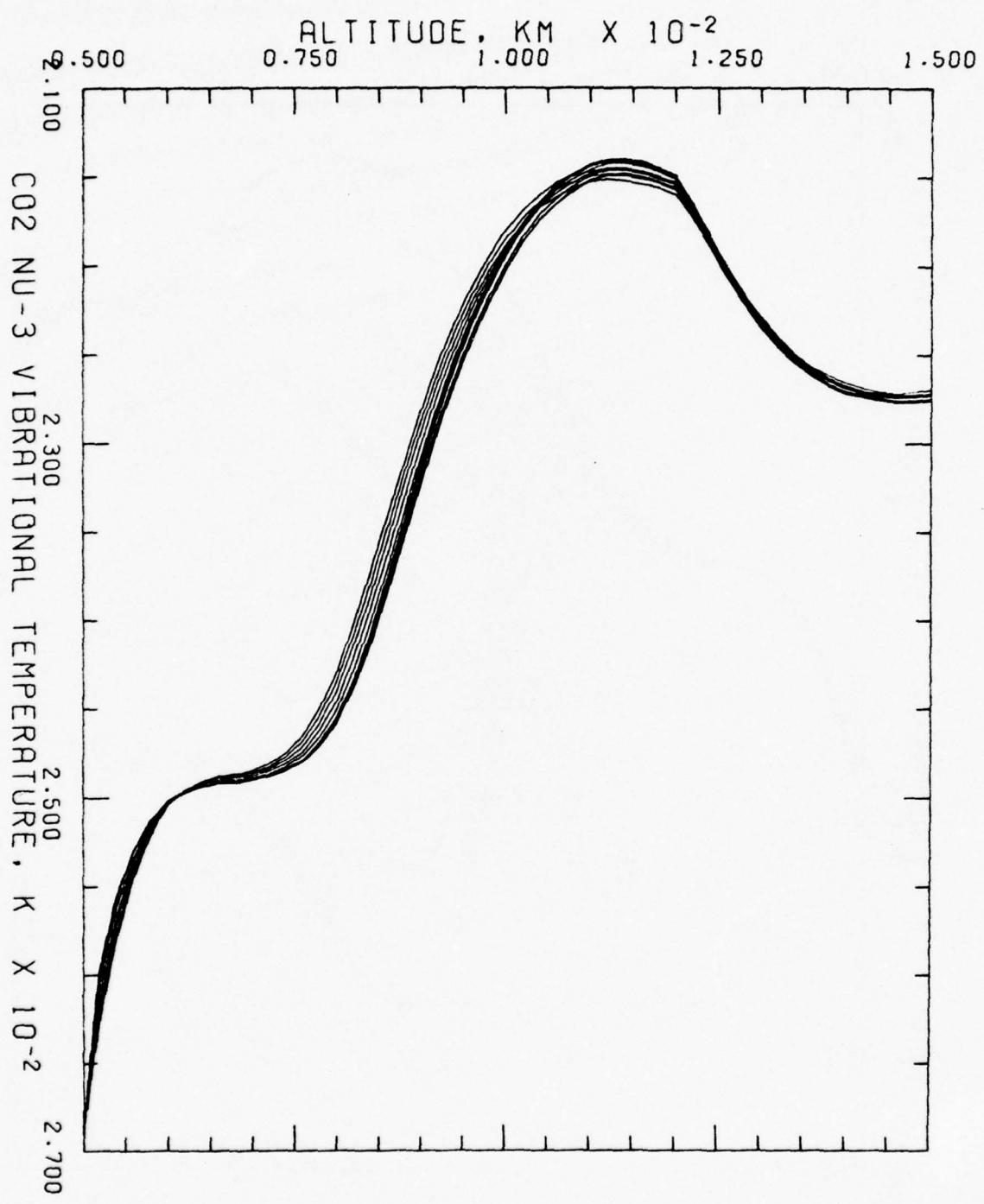


Figure 10

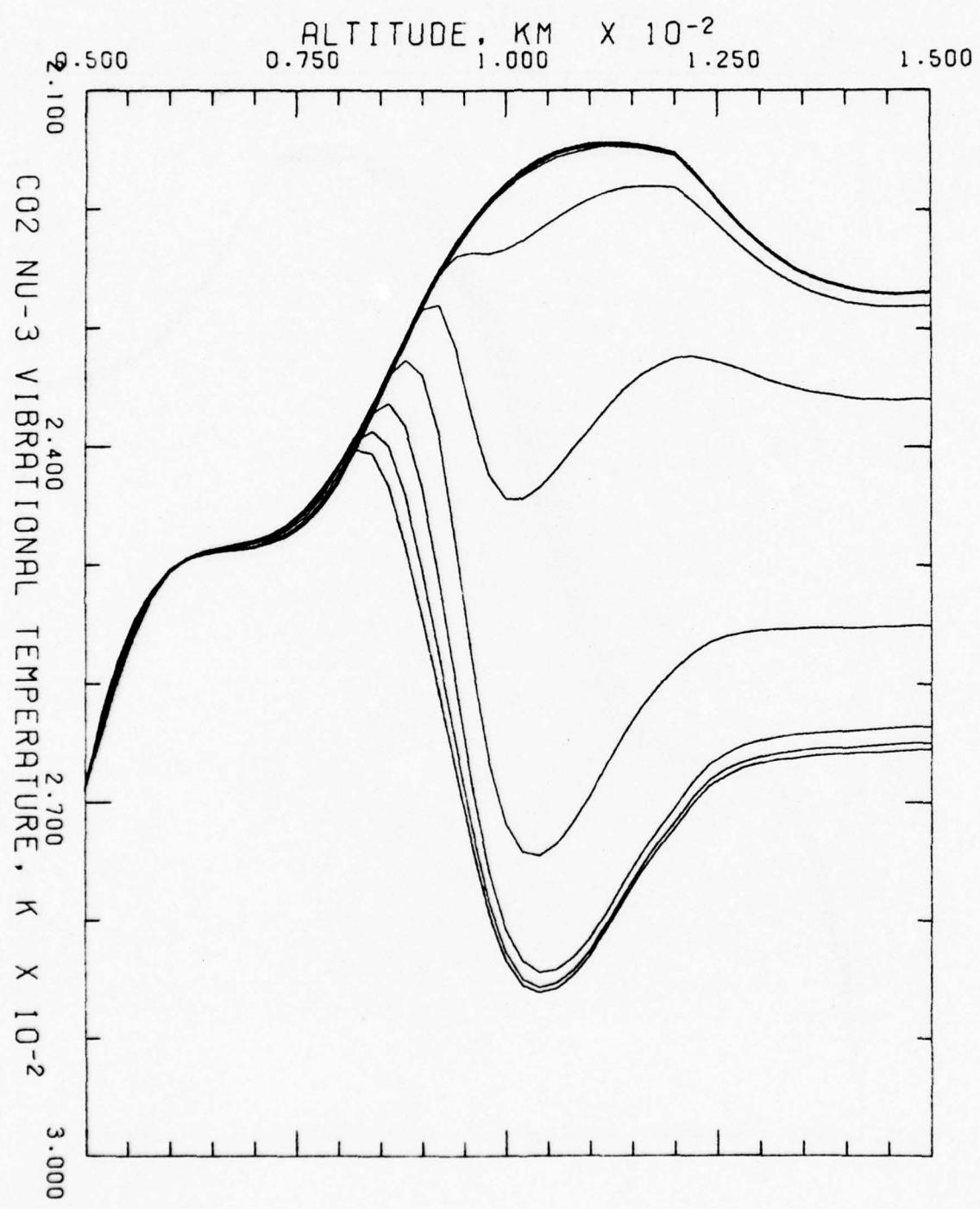


Figure 11

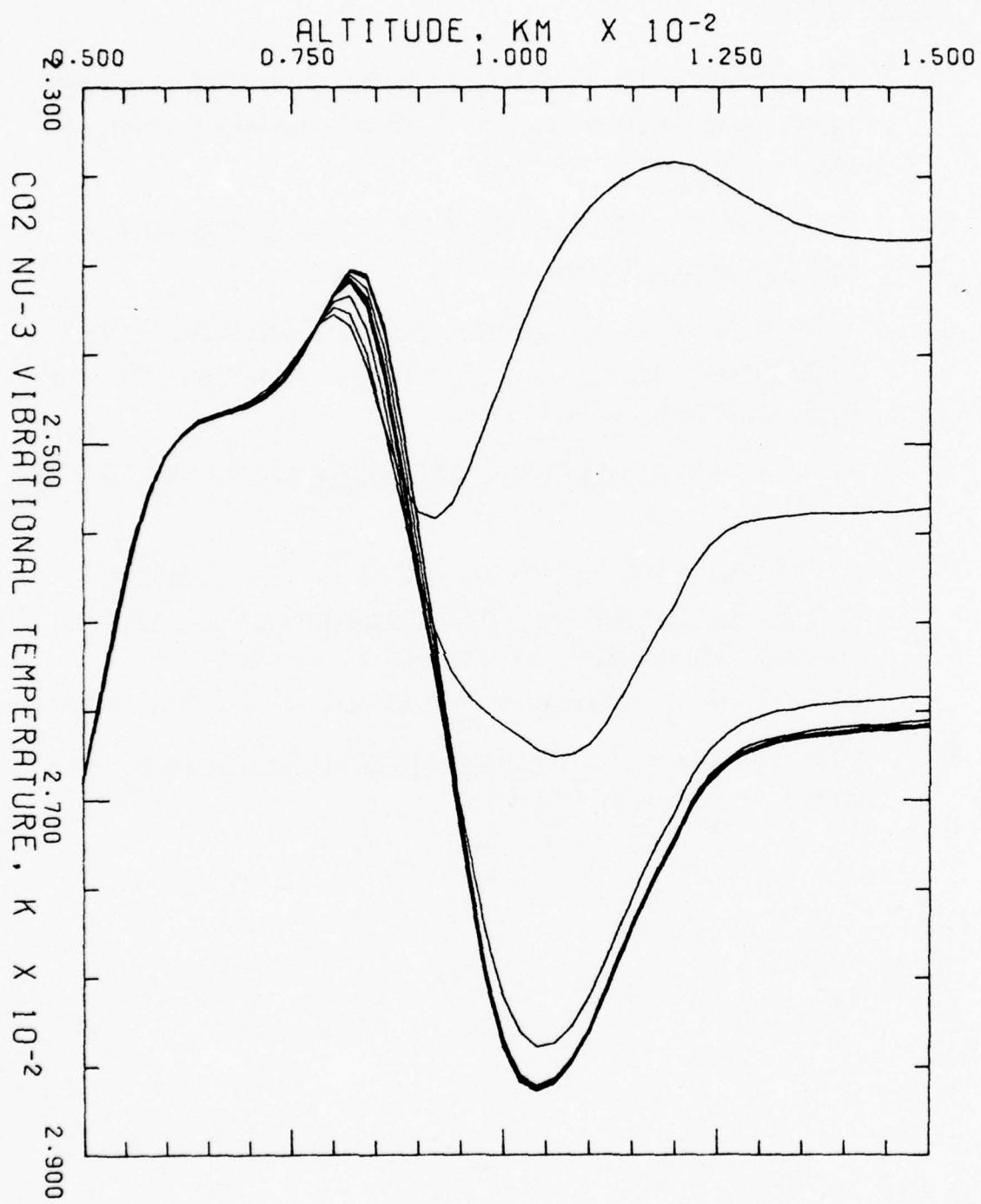


Figure 12

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1. M. J. Berger and S. M. Seltzer, "Energy Deposition by Auroral Electrons in the Atmosphere", J. of Atmospheric Terrestrial Physics, 32, 1970.
2. T. C. Degges, AFCRL-TR-74-0606, A High Altitude Infrared Radiance Model, December 1974.
3. V. L. Corbin et al, Atmospheric Radiance Models for Limb-Viewing Geometry in the Five- to Twenty-Five Micron Spectral Region, AFCRL-69-0552 (1969).
4. T. C. Degges, A High Altitude Radiance Model, AFCRL 72-0273 (1972).
5. Garvin, D., and Hampson, R. F., Editors, Chemical Kinetics Data Survey VII. Tables of Rate and Photochemical Data for Modelling of the Stratosphere (Revised). NBSIR 74-430, National Bureau of Standards, Washington, D. C., January 1974.
6. O. P. Manley et al, Modelling of Optical IR Backgrounds, AFCRL-TR-73-0565, (1973).

Appendix 1

Auroral Electron Energy Spectra

Rocket A18.219-1

Launched 25 February 1974

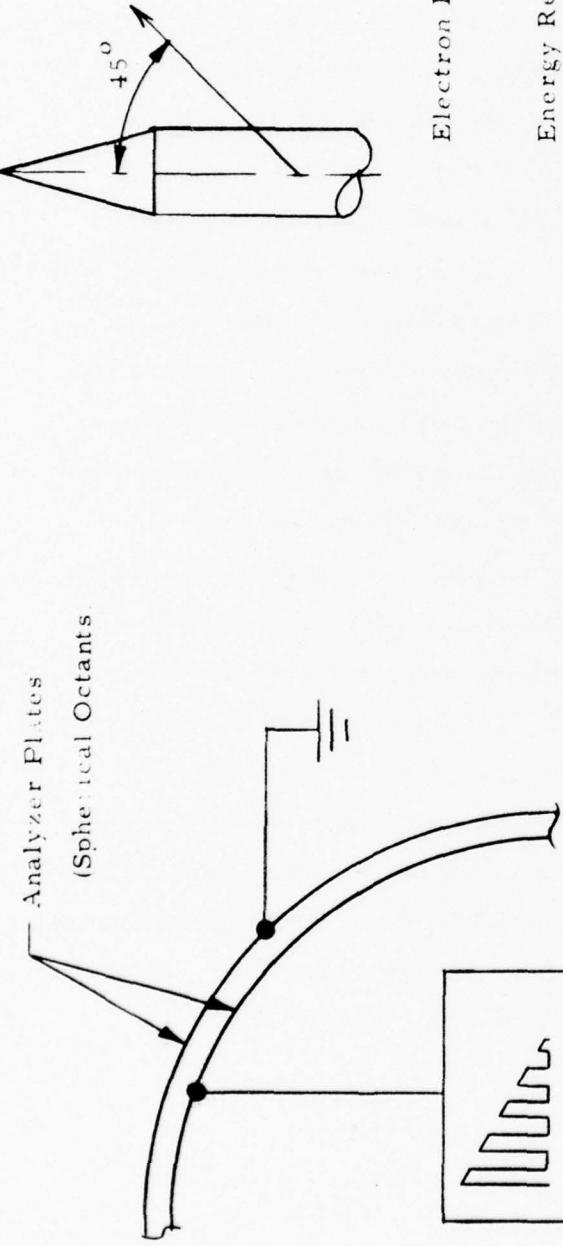
The data in the following figures was presented at the DNA ICECAP Data Meeting in January 1975. Rocket A 18.219-1 was launched 25 February 1975 and overflew the same bright auroral arc twice. Peak energy deposition was in excess of $100 \text{ ergs cm}^{-2} \text{ sec}^{-1}$. The electron energy deposition was measured with two instruments, an aluminum covered scintillator coupled to a photomultiplier and an electrostatic analyzer (ESA). Figure 1 shows schematically the operation of the ESA. Spherical analyzer plates are used to select a narrow energy band. The electrons are then accelerated with a voltage applied to a thin aluminum deposition on the face of a calcium fluoride scintillator. This extends the low energy response of the instrument. By modulating the analyzer plate voltage and using a narrow band amplifier the sensitivity of the instrument is significantly improved. Figure 2 shows the scintillator measurement as a function of time. This detector measures the integrated electron flux above 4 or 5 keV. The two peaks correspond to the time when the rocket is on the magnetic field line connecting to the visible aurora. The auroral motion is presently being analyzed as part of the data analysis. The visible aurora initially is to the north, moves to the south under the rocket, and then moves to the north again passing under the rocket.

The following figures show the auroral energy spectrum measured with the ESA. The insert shows the scintillator measurement and the inverted triangle indicates when that electron energy spectrum was taken. The auroral peak energy extended beyond 30 keV and showed significant spectral differences north and south of the peak region. The spectral form north of the peak region was much softer than that south of the peak region. These spectral differences persisted throughout the flight. The most striking feature is that the electron spectral description of the auroral form varied very little while the form moved rapidly in the sky. The detailed analysis of this aurora will be presented in a later report.

The data in the following figures was presented at the DNA ICECAP Data Meeting in January 1975. Rocket A 18.219-1 was launched 25 February 1975 and overflew the same bright auroral arc twice. Peak energy deposition was in excess of $100 \text{ ergs cm}^{-2} \text{ sec}^{-1}$. The electron energy deposition was measured with two instruments, an aluminum covered scintillator coupled to a photomultiplier and an electrostatic analyzer (ESA). Figure 1 shows schematically the operation of the ESA. Spherical analyzer plates are used to select a narrow energy band. The electrons are then accelerated with a voltage applied to a thin aluminum deposition on the face of a calcium fluoride scintillator. This extends the low energy response of the instrument. By modulating the analyzer plate voltage and using a narrow band amplifier the sensitivity of the instrument is significantly improved. Figure 2 shows the scintillator measurement as a function of time. This detector measures the integrated electron flux above 4 or 5 keV. The two peaks correspond to the time when the rocket is on the magnetic field line connecting to the visible aurora. The auroral motion is presently being analyzed as part of the data analysis. The visible aurora initially is to the north, moves to the south under the rocket, and then moves to the north again passing under the rocket.

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Analyzer Plates
(Spherical Octants)

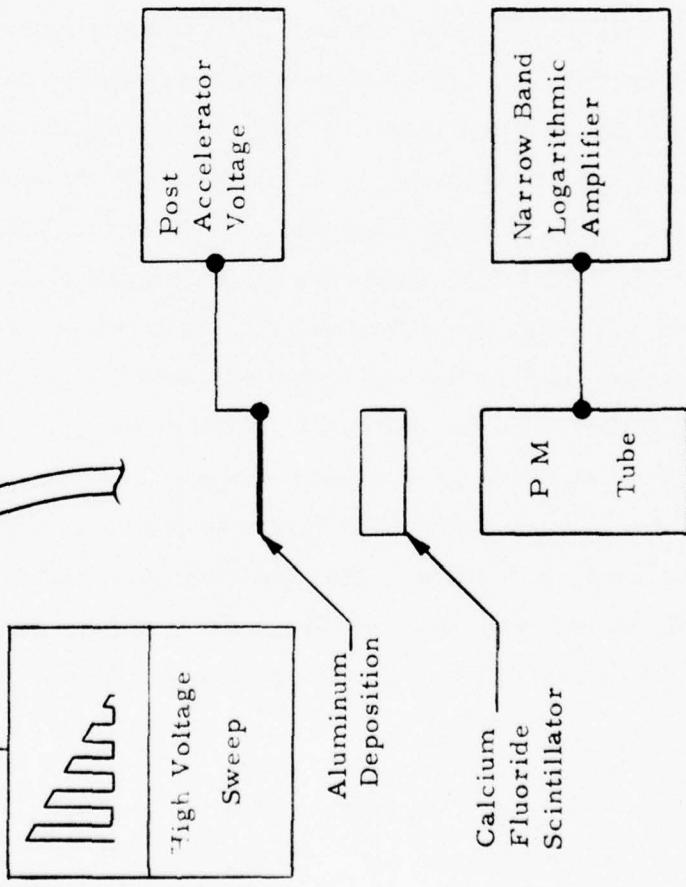


Electron Energy - 3 kev to
30 kev

Energy Resolution - $11\sigma_0$

Angular Field of View -
 $6.4^\circ \times 10^\circ$

Geometric Factor -
 $9.2 \times 10^{-2} \text{ cm}^2 \text{ ster}$



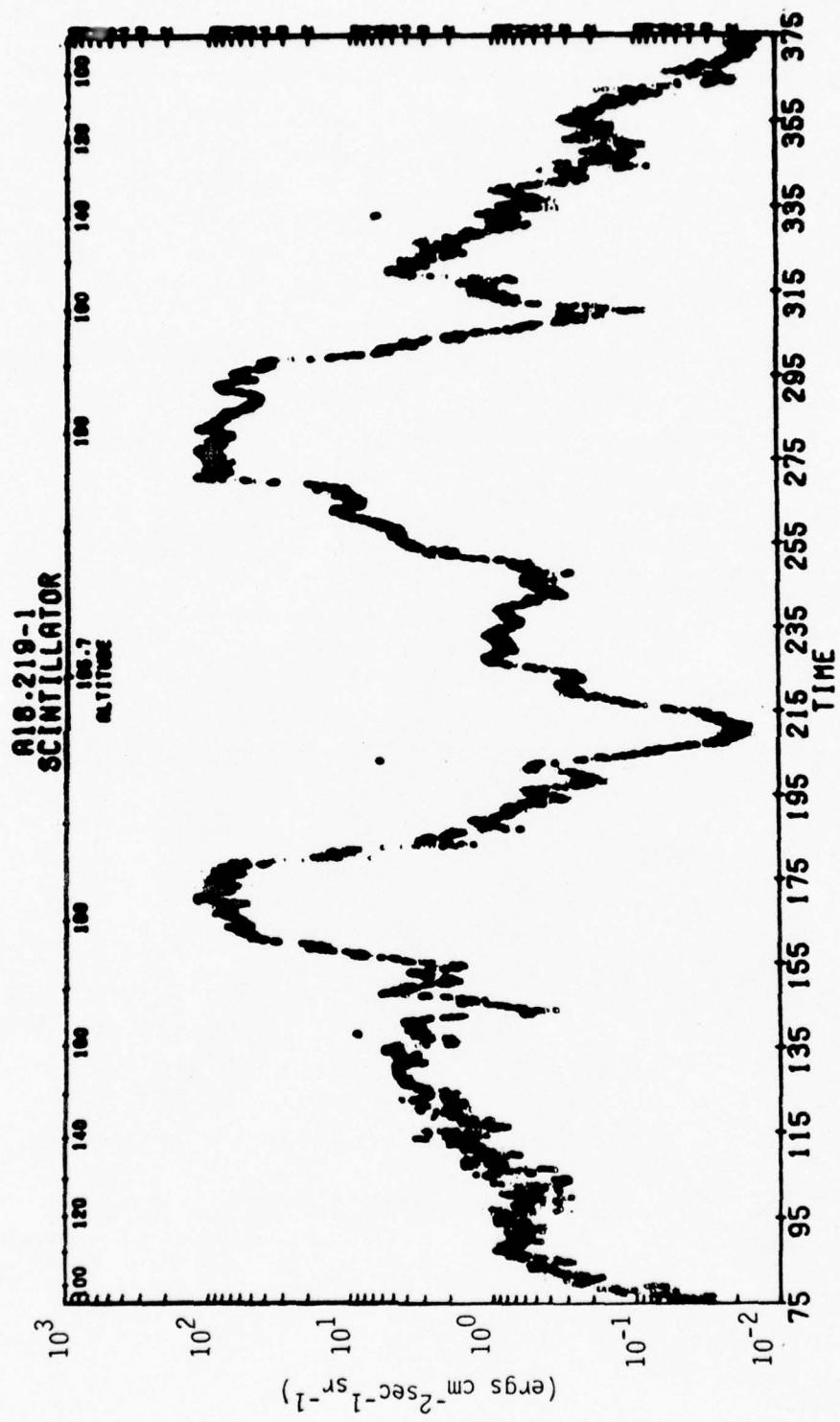


FIGURE 2:

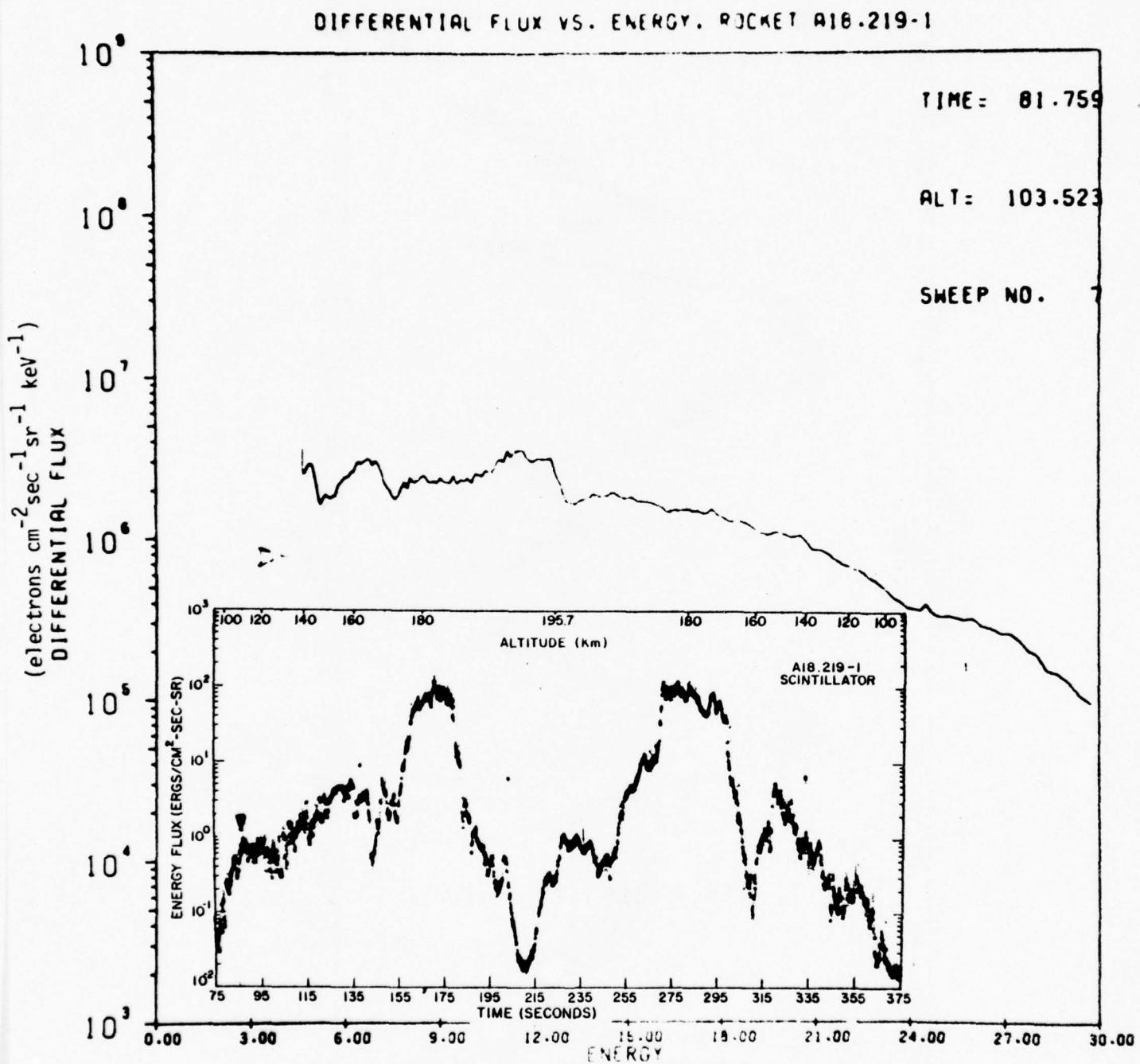


FIGURE 3:

DIFFERENTIAL FLUX VS. ENERGY. ROCKET A18.219-1

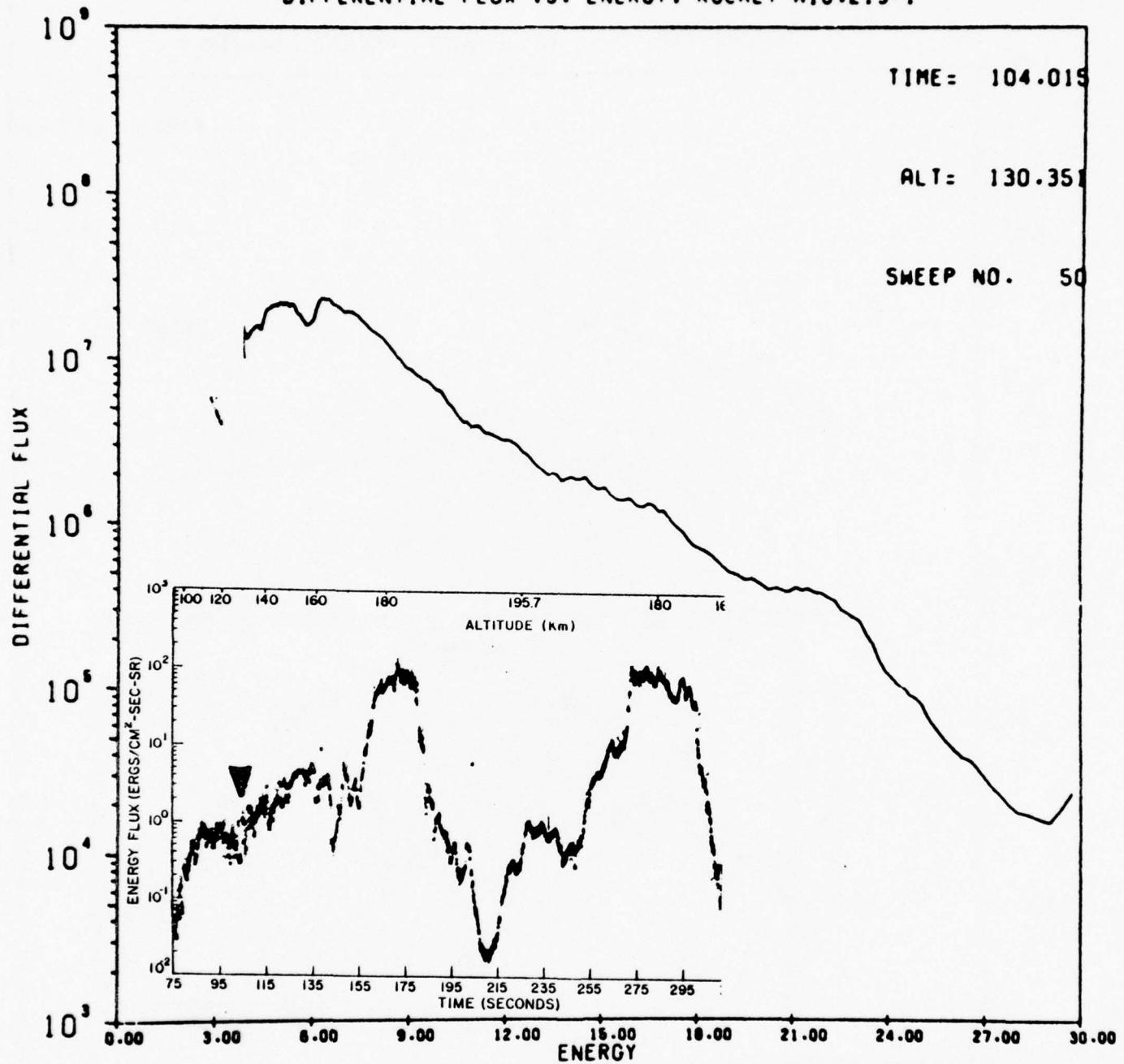


FIGURE 4:

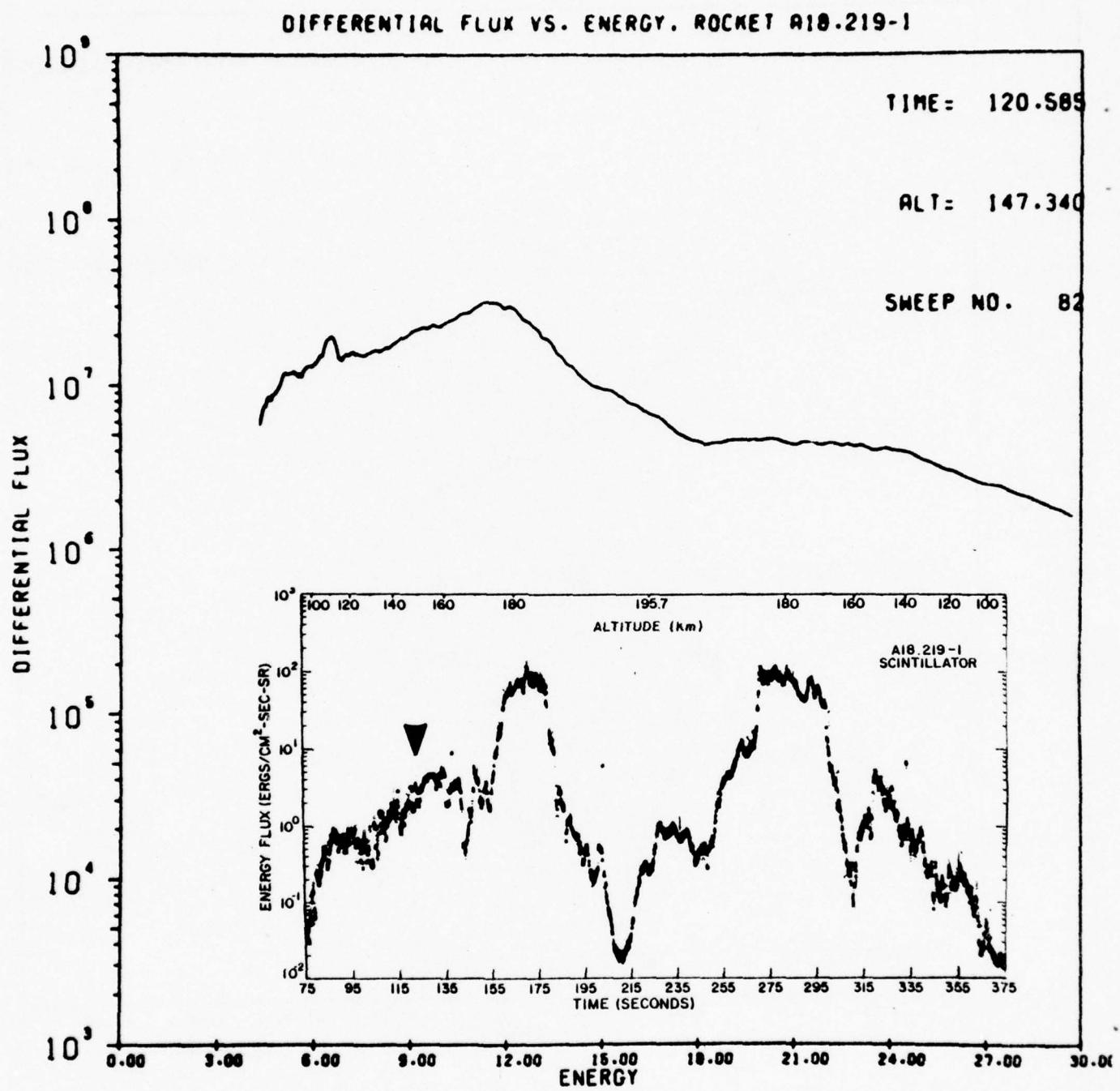


FIGURE 5:

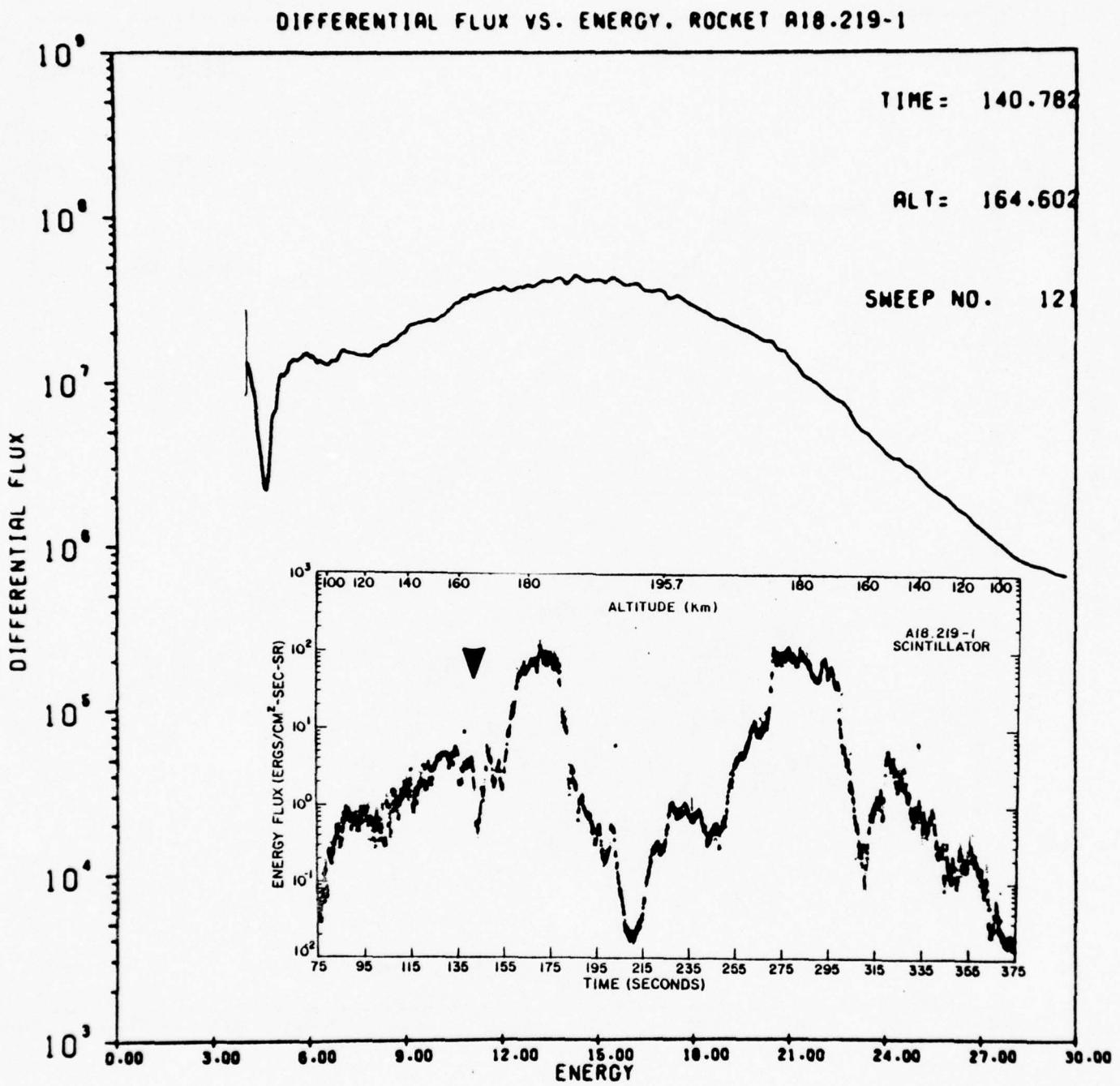


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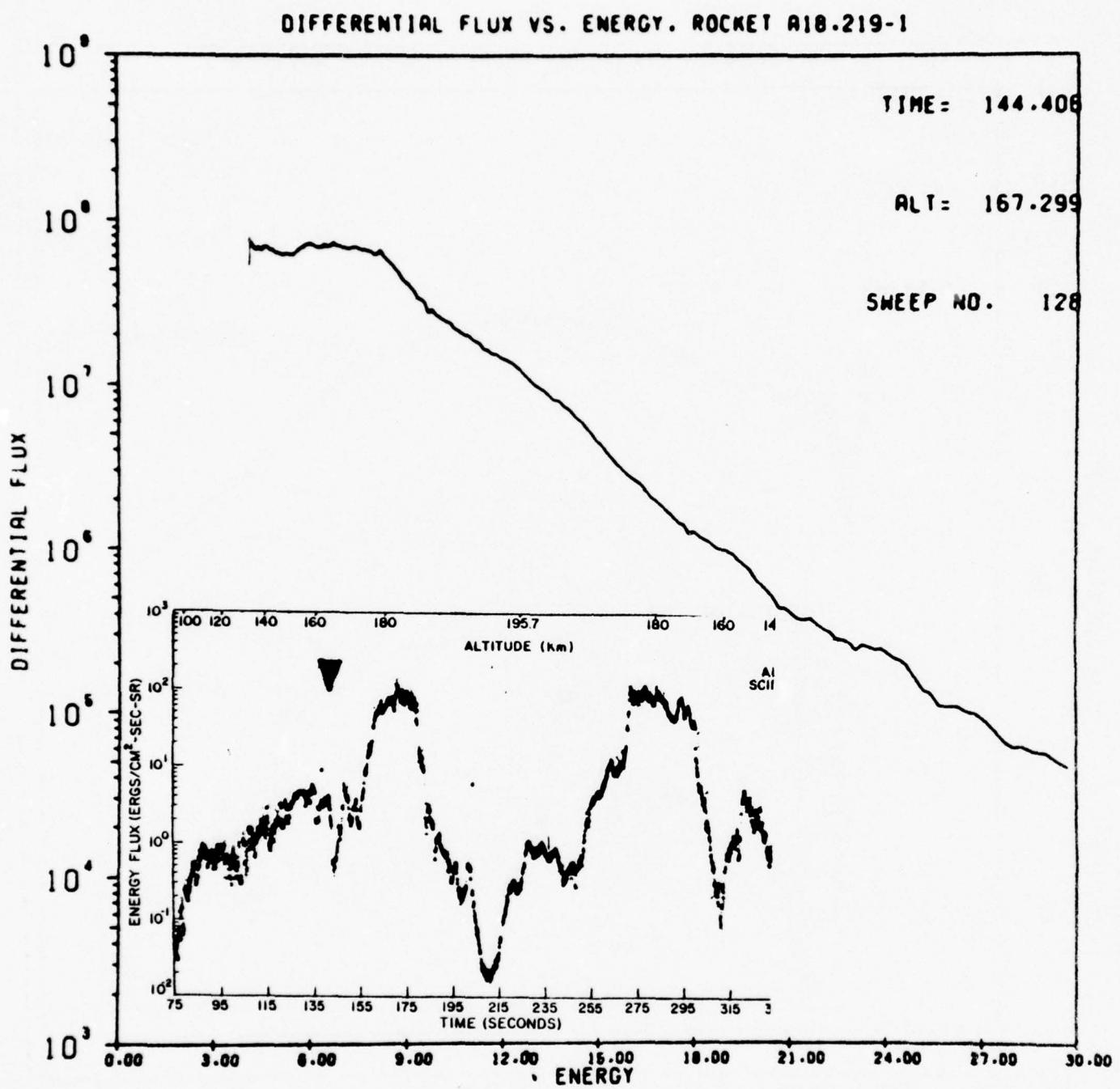


FIGURE 7:

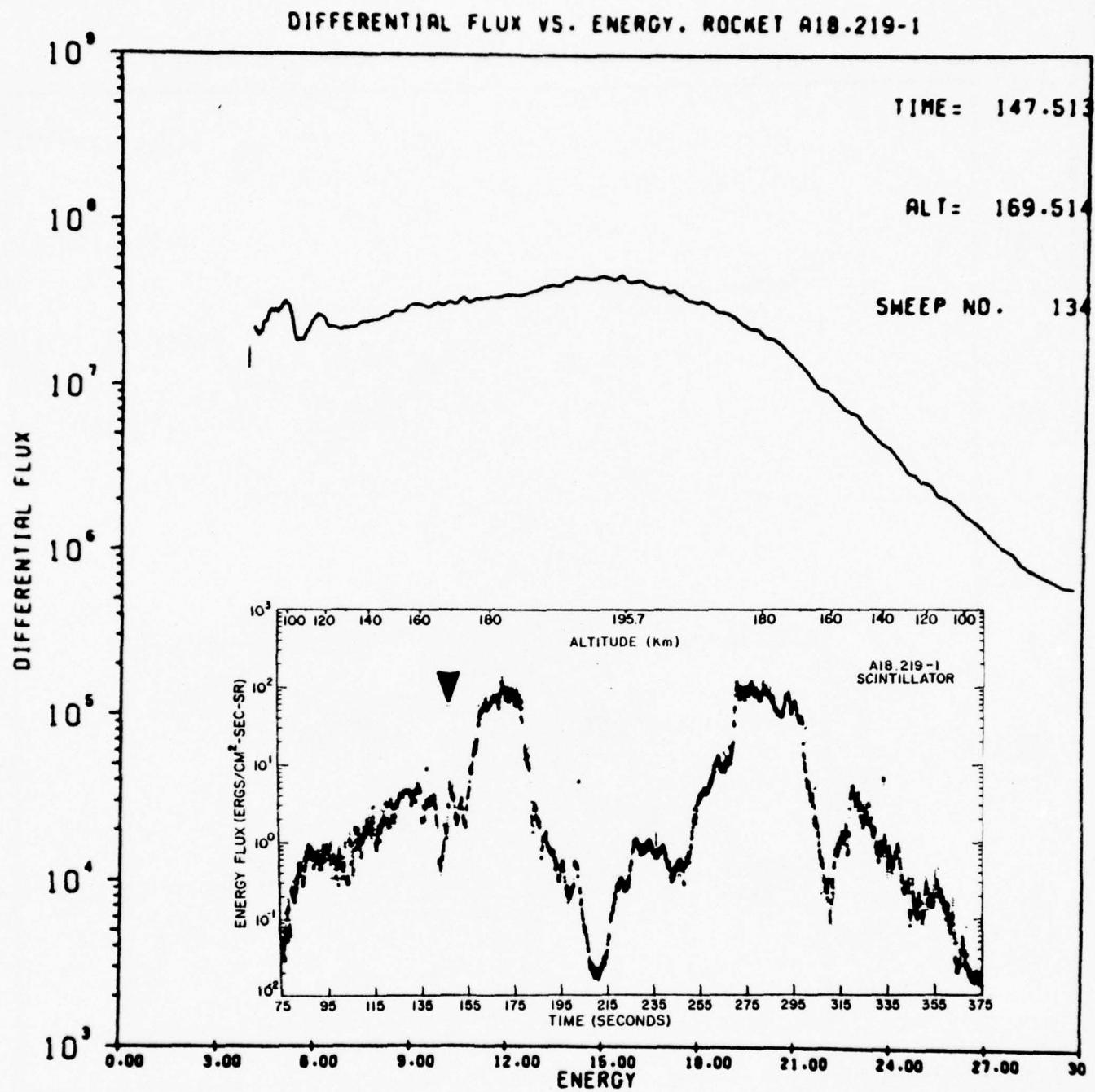


FIGURE 8:

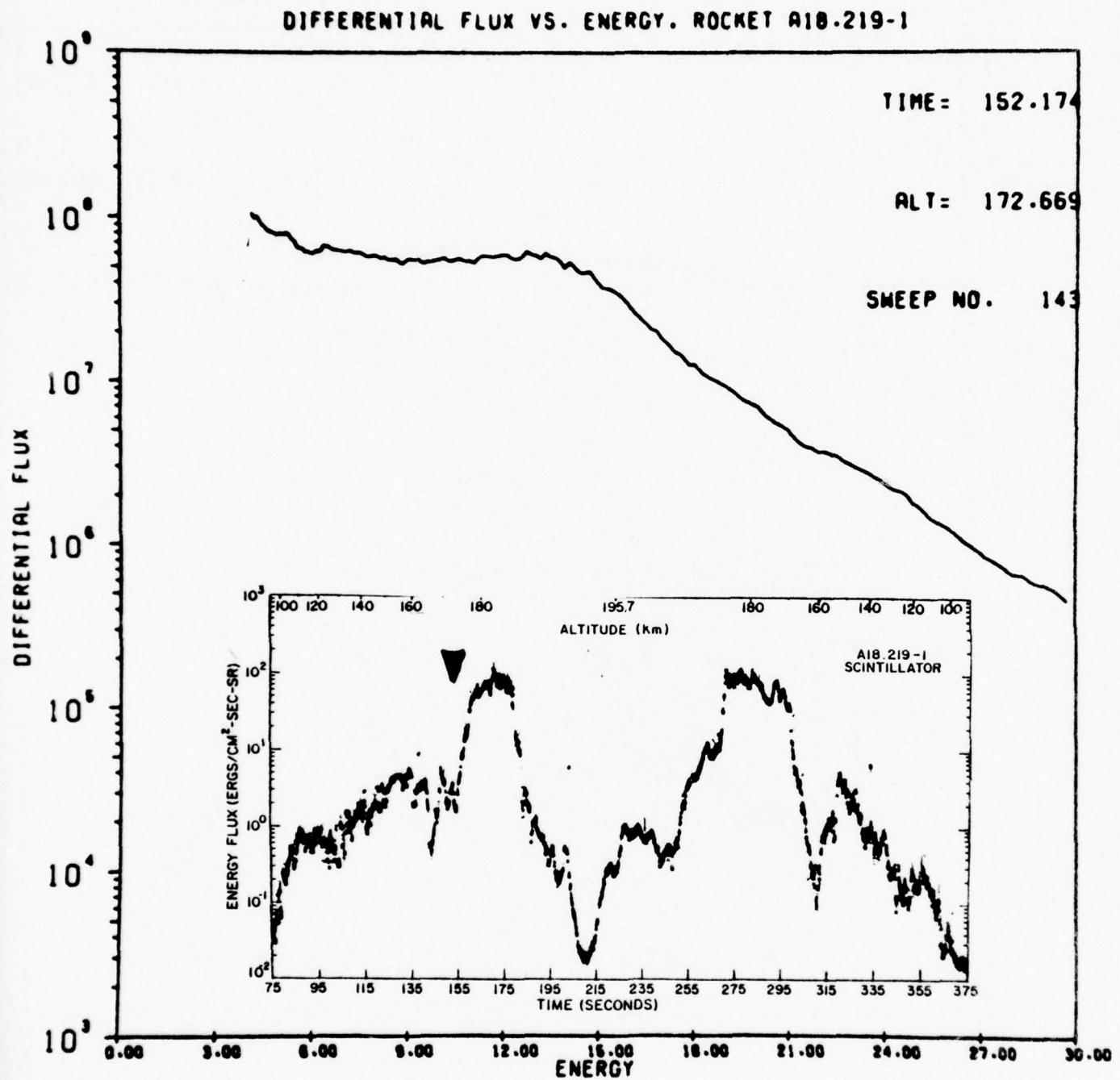


FIGURE 9:

DIFFERENTIAL FLUX VS. ENERGY. ROCKET A18.219-1

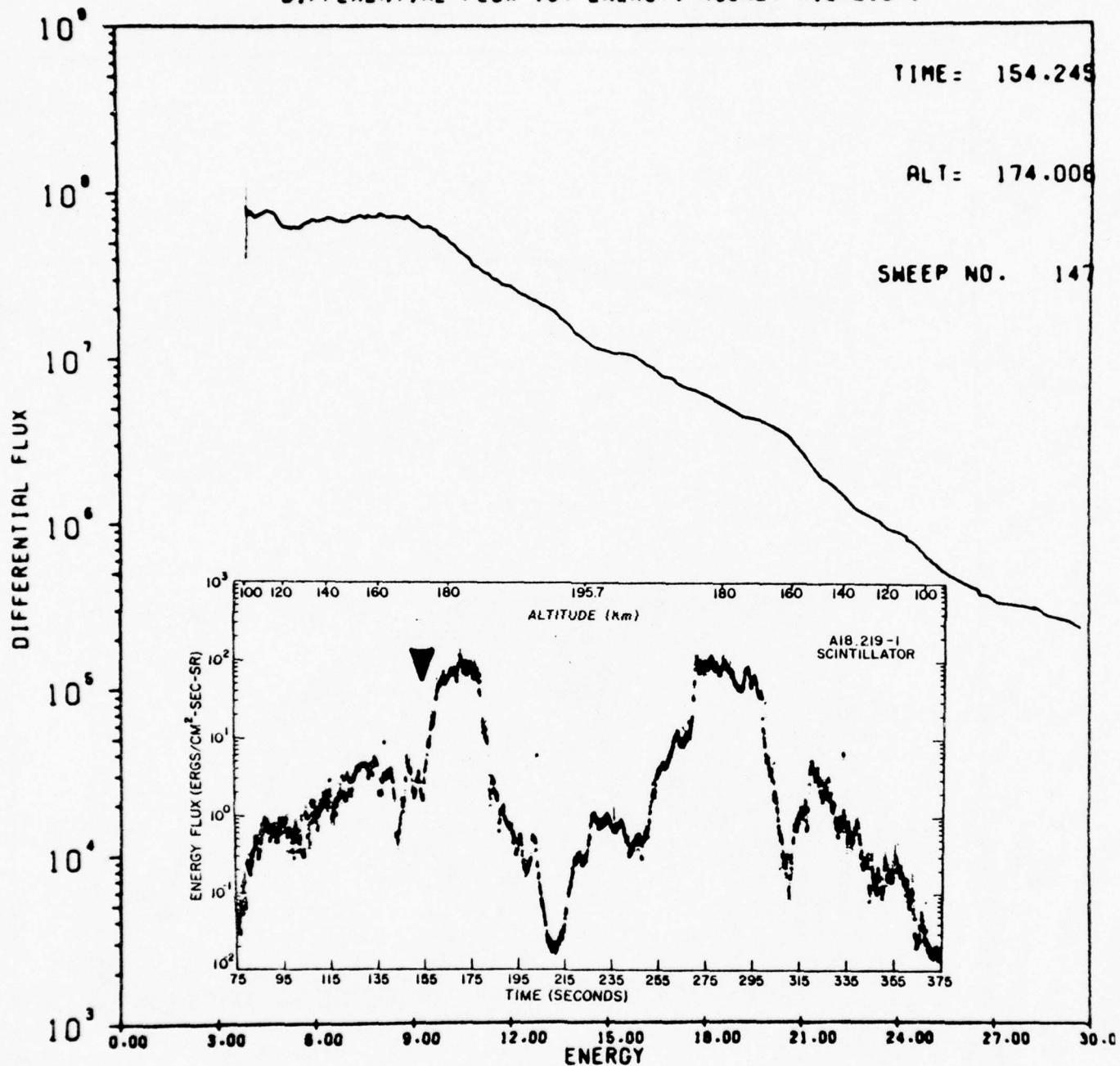


FIGURE 10:

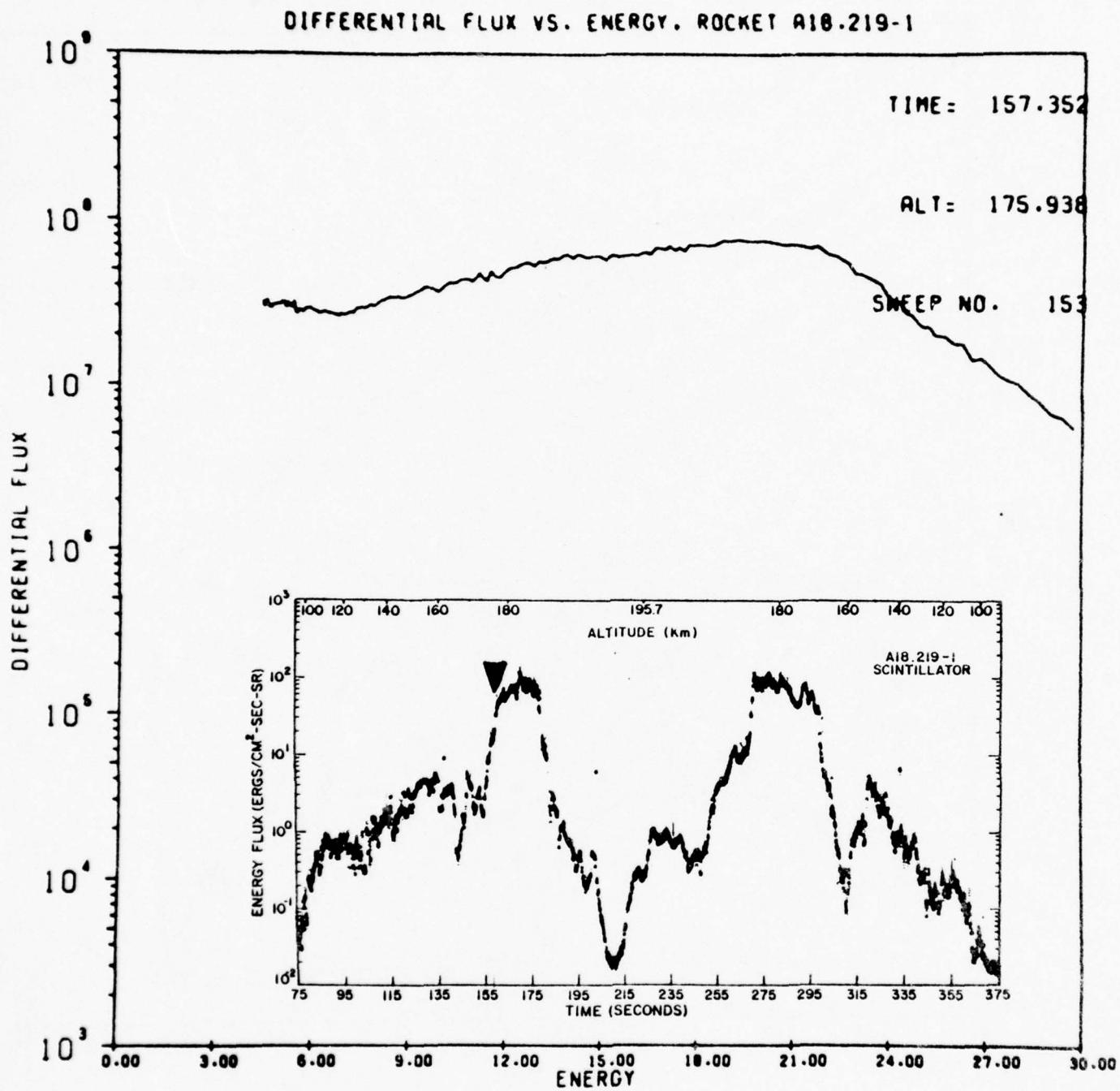


FIGURE 11:

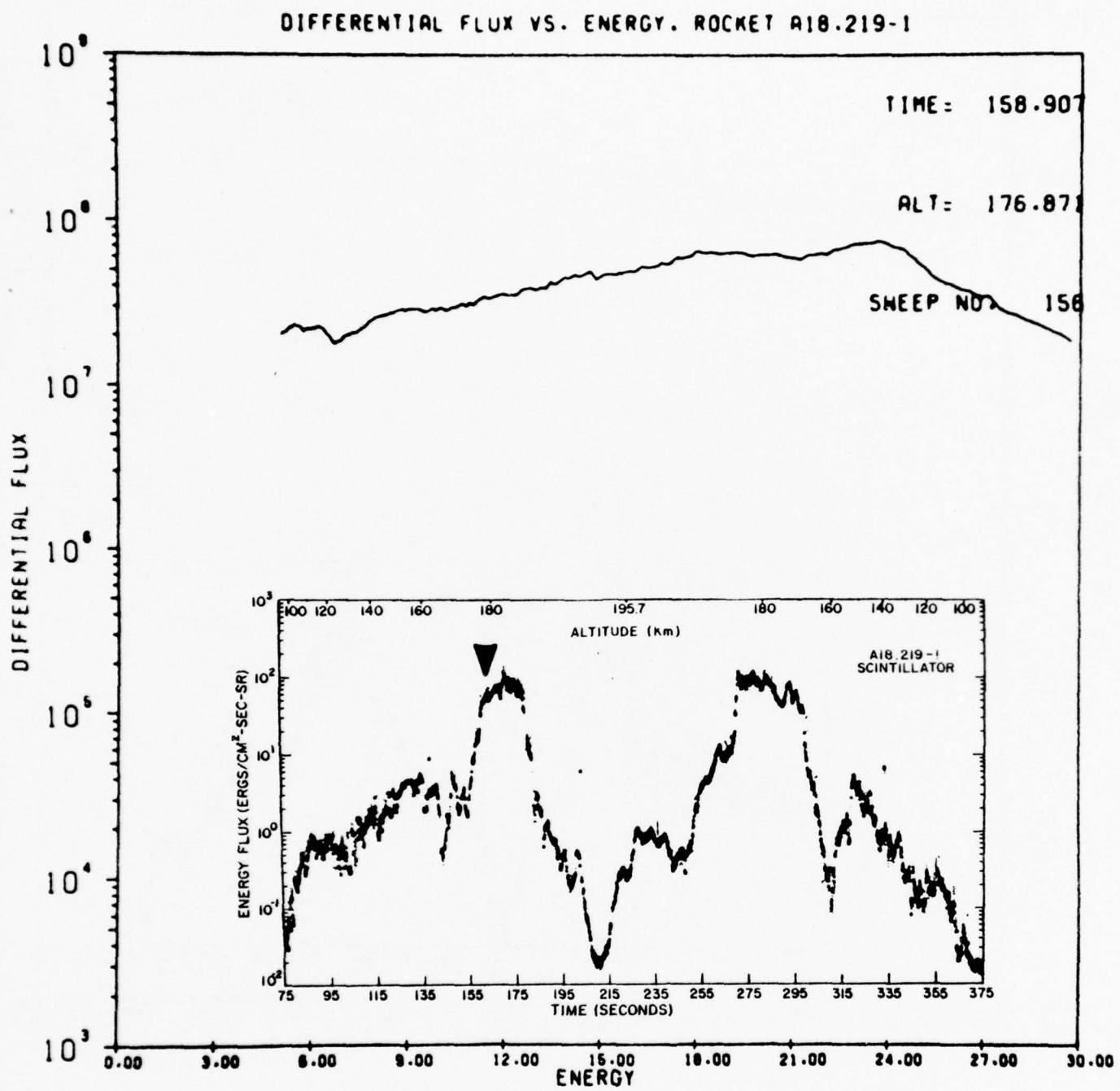


FIGURE 12:

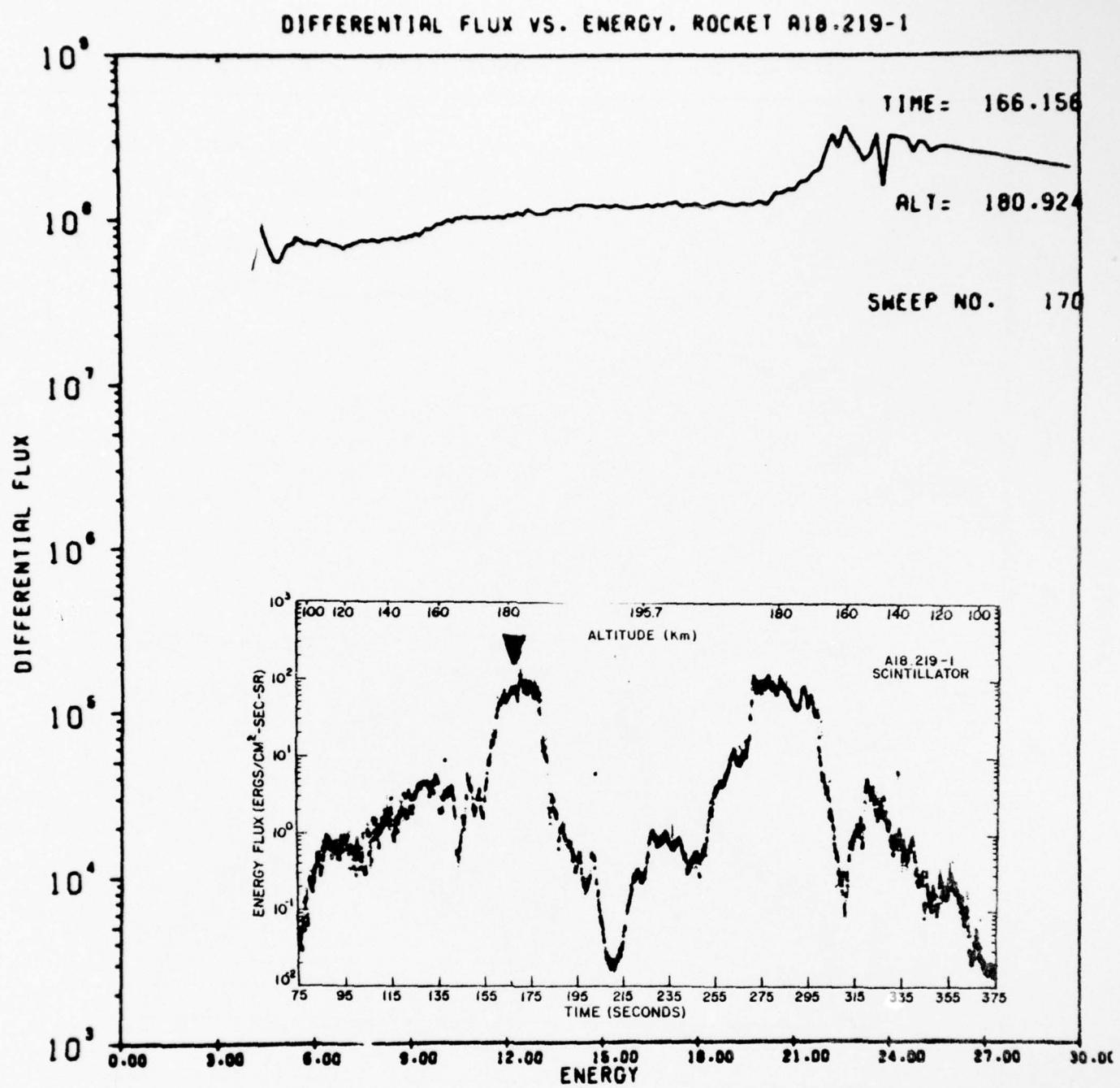


FIGURE 13

DIFFERENTIAL FLUX VS. ENERGY. ROCKET A18.219-1

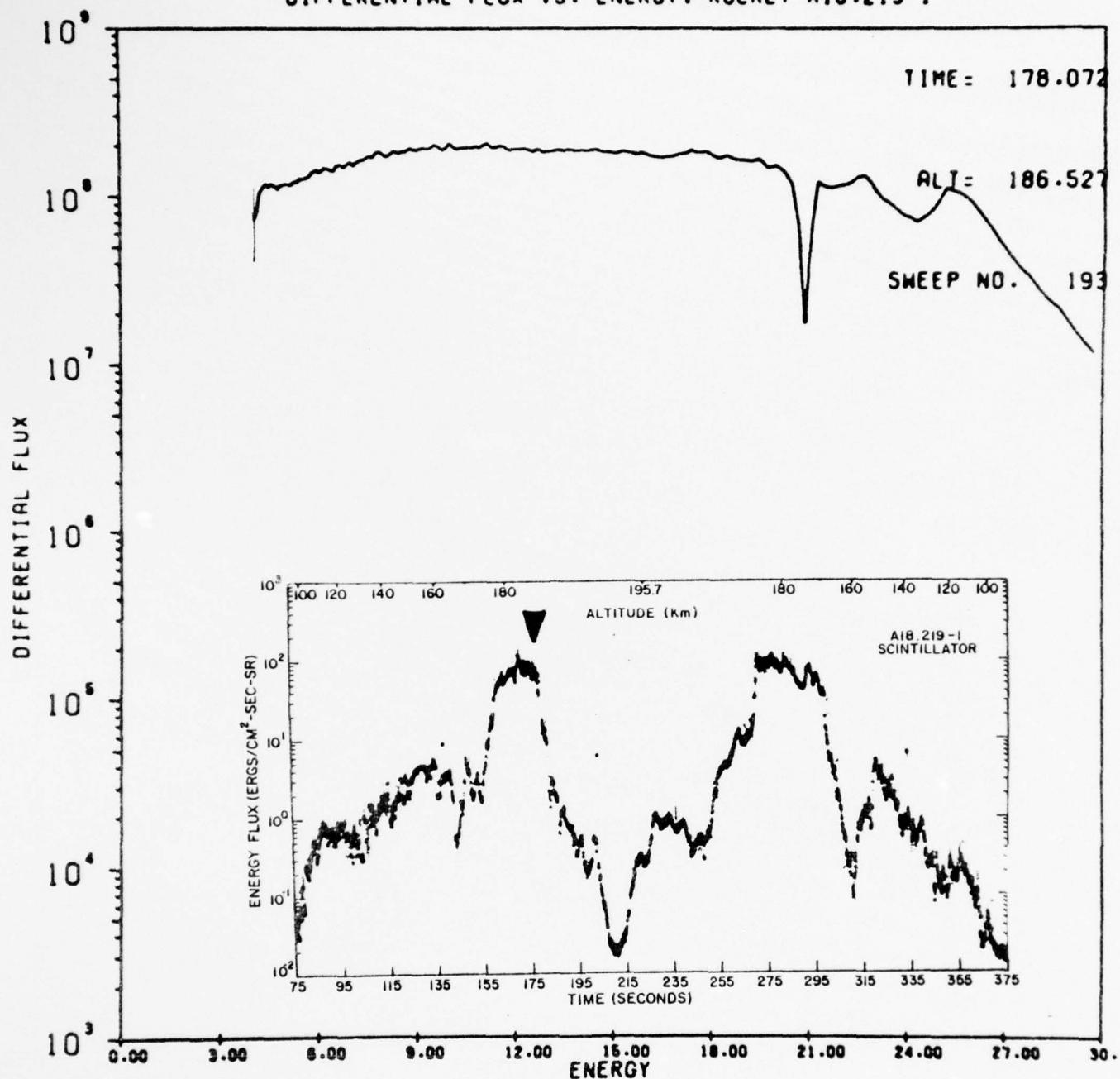


FIGURE 14

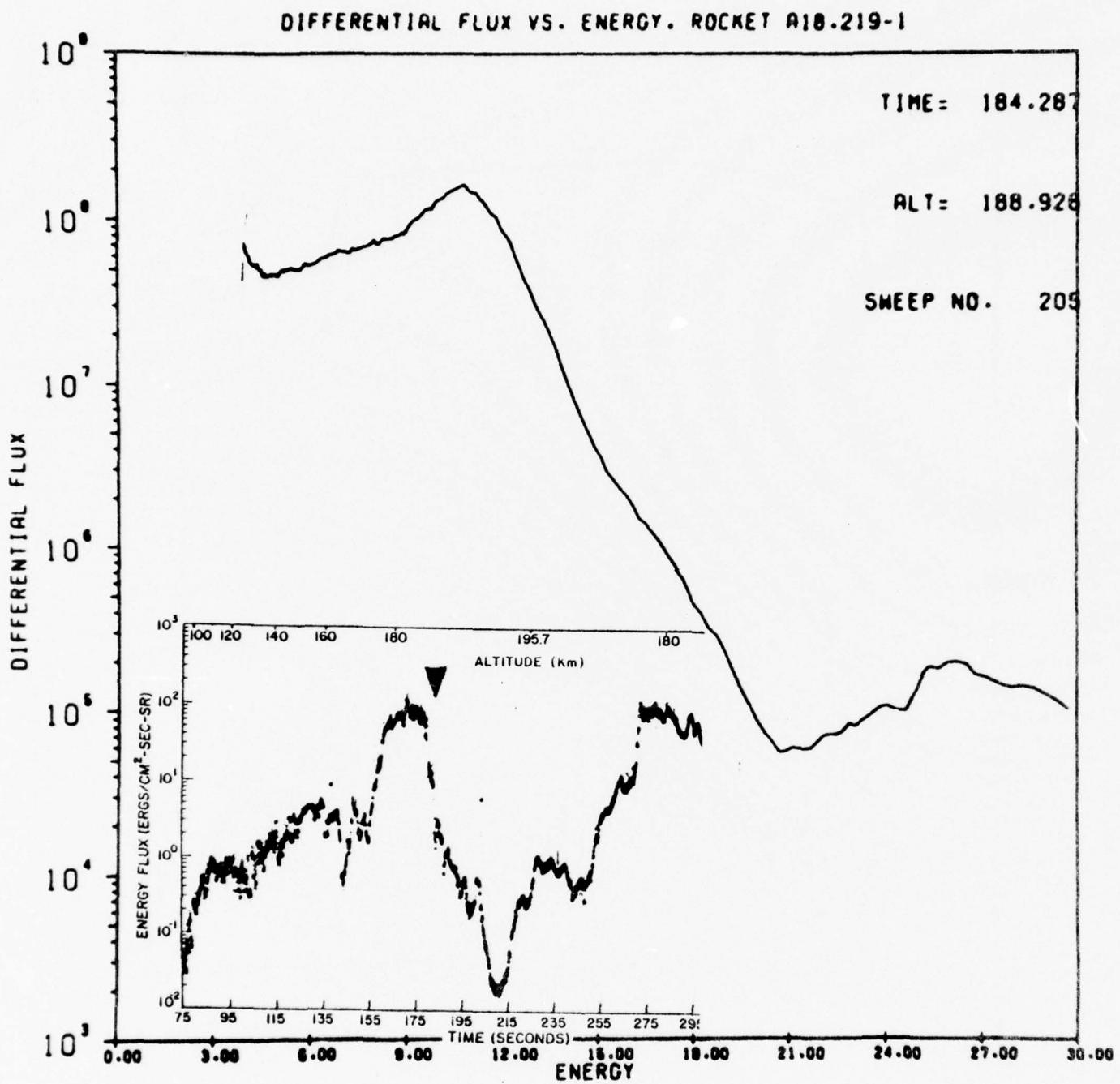


FIGURE 15

DIFFERENTIAL FLUX VS. ENERGY. ROCKET A18.219-1

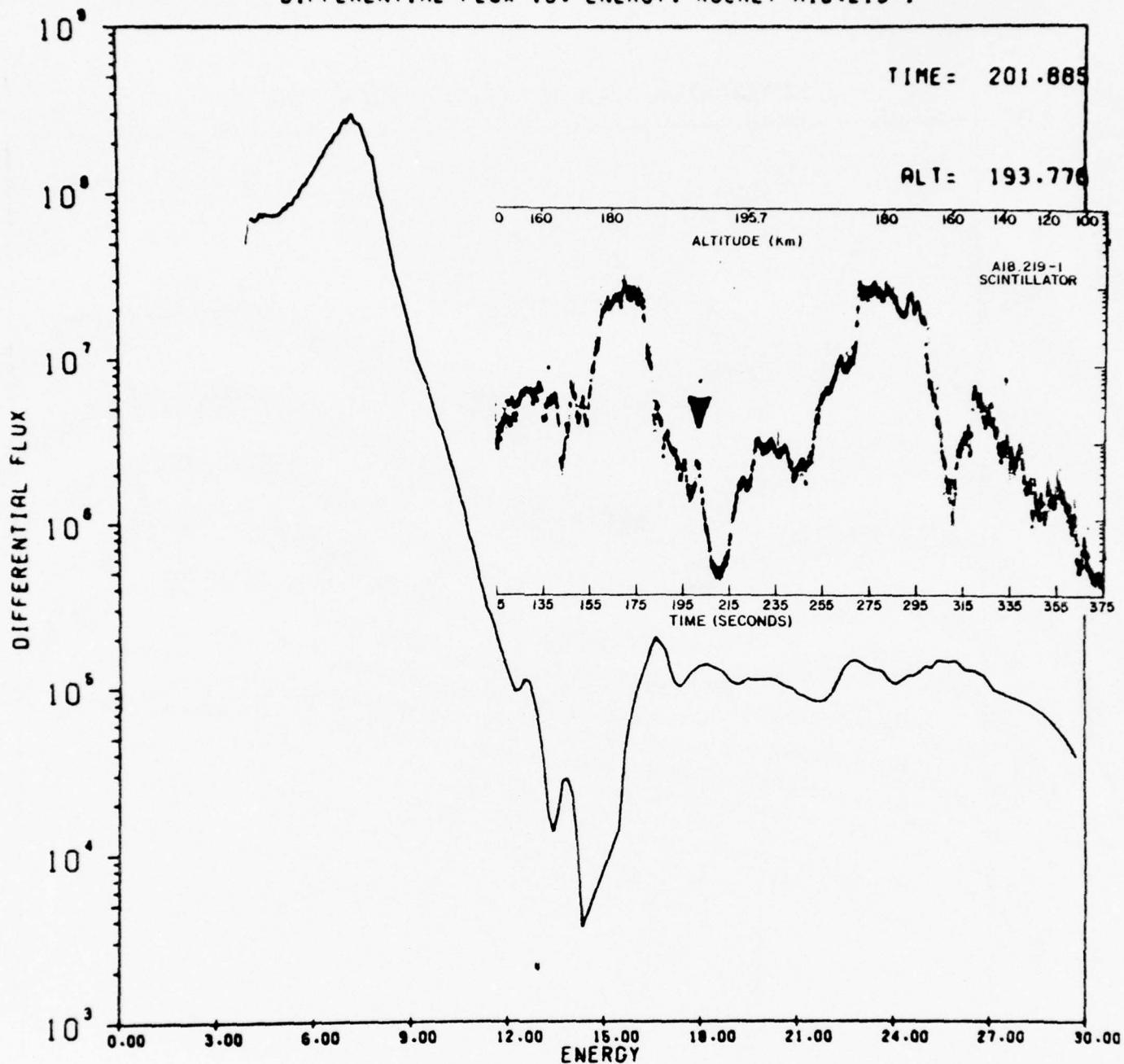


FIGURE 16

DIFFERENTIAL FLUX VS. ENERGY. ROCKET A10.219-1

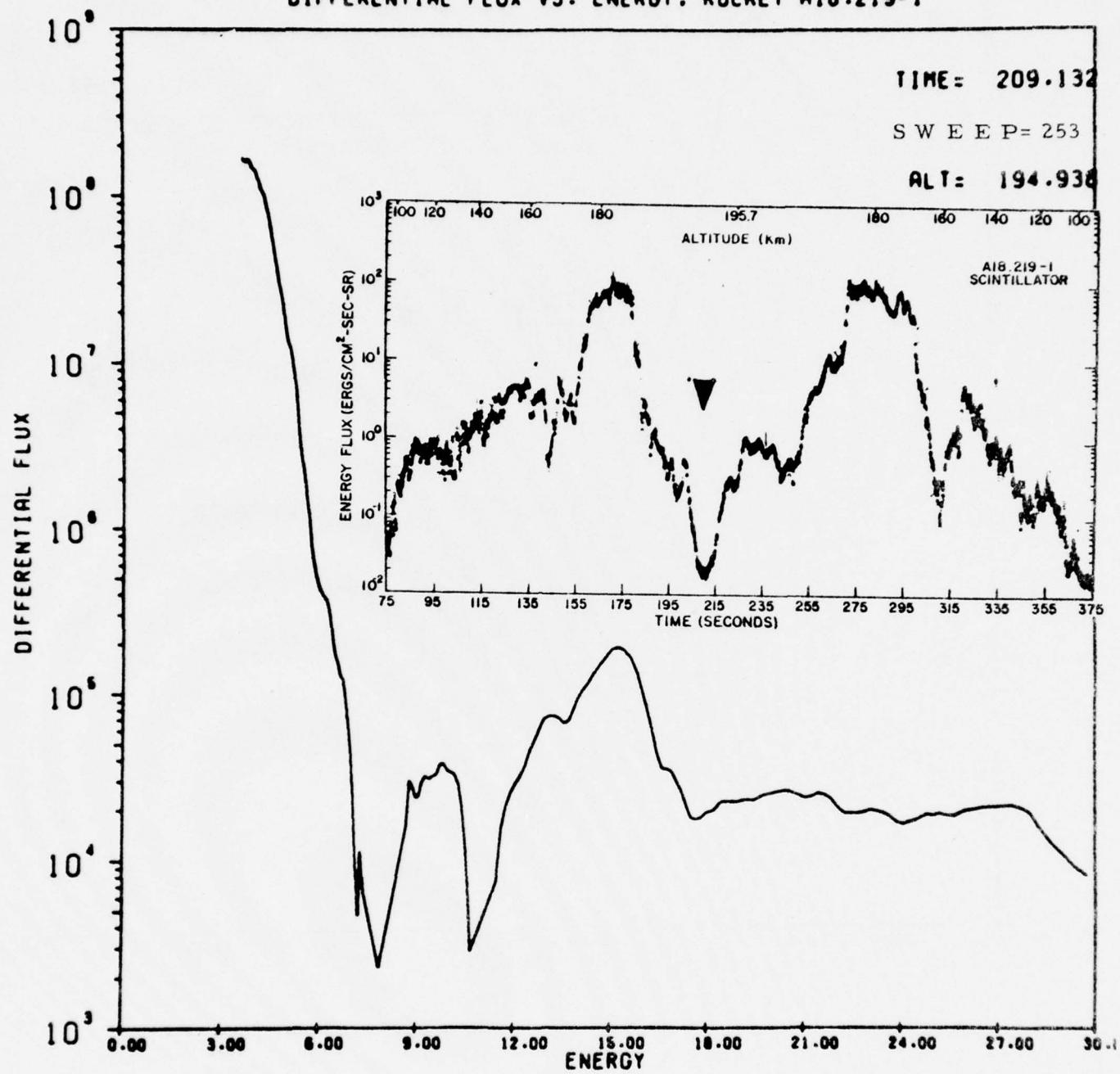


FIGURE 17

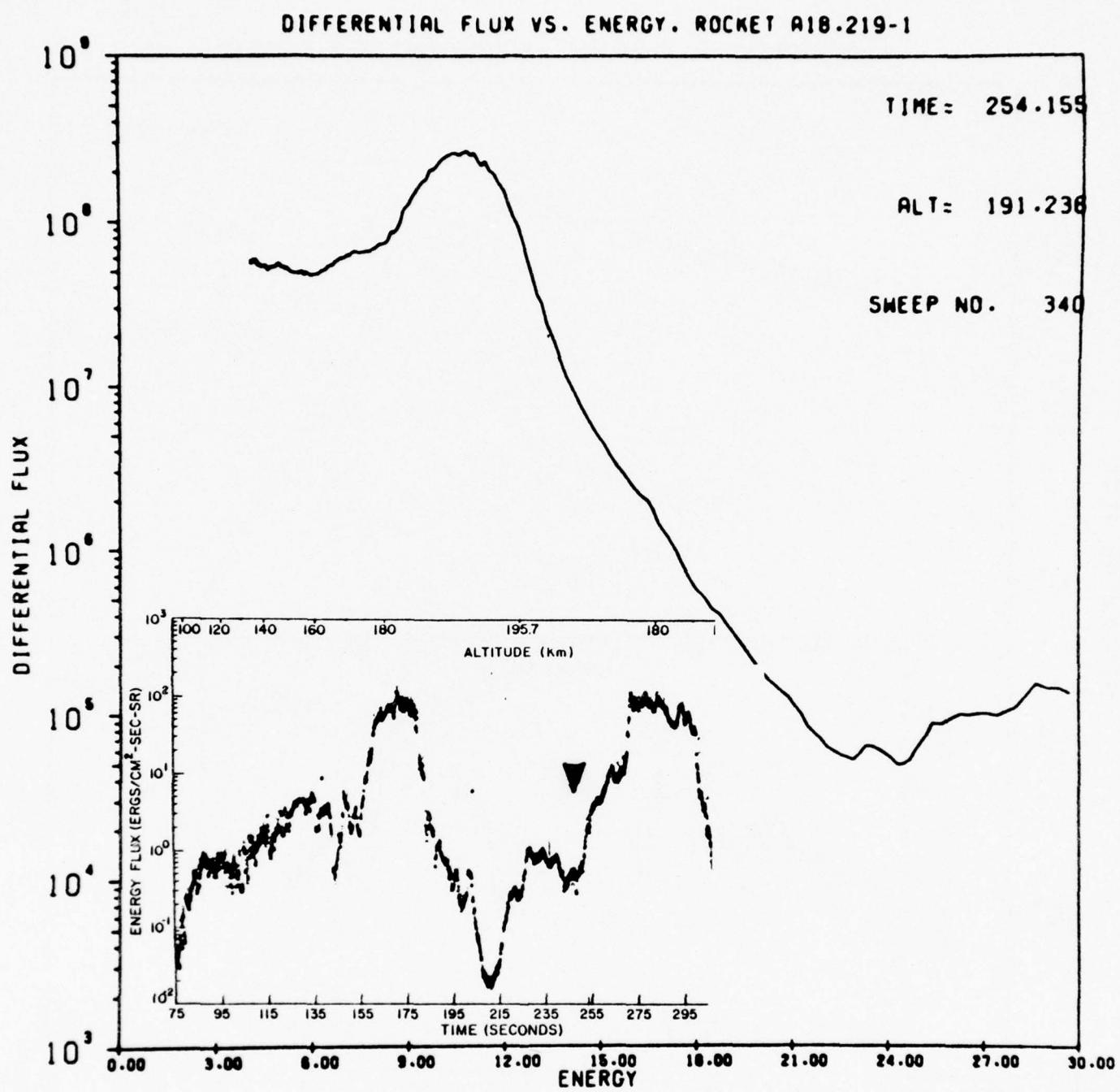


FIGURE 18

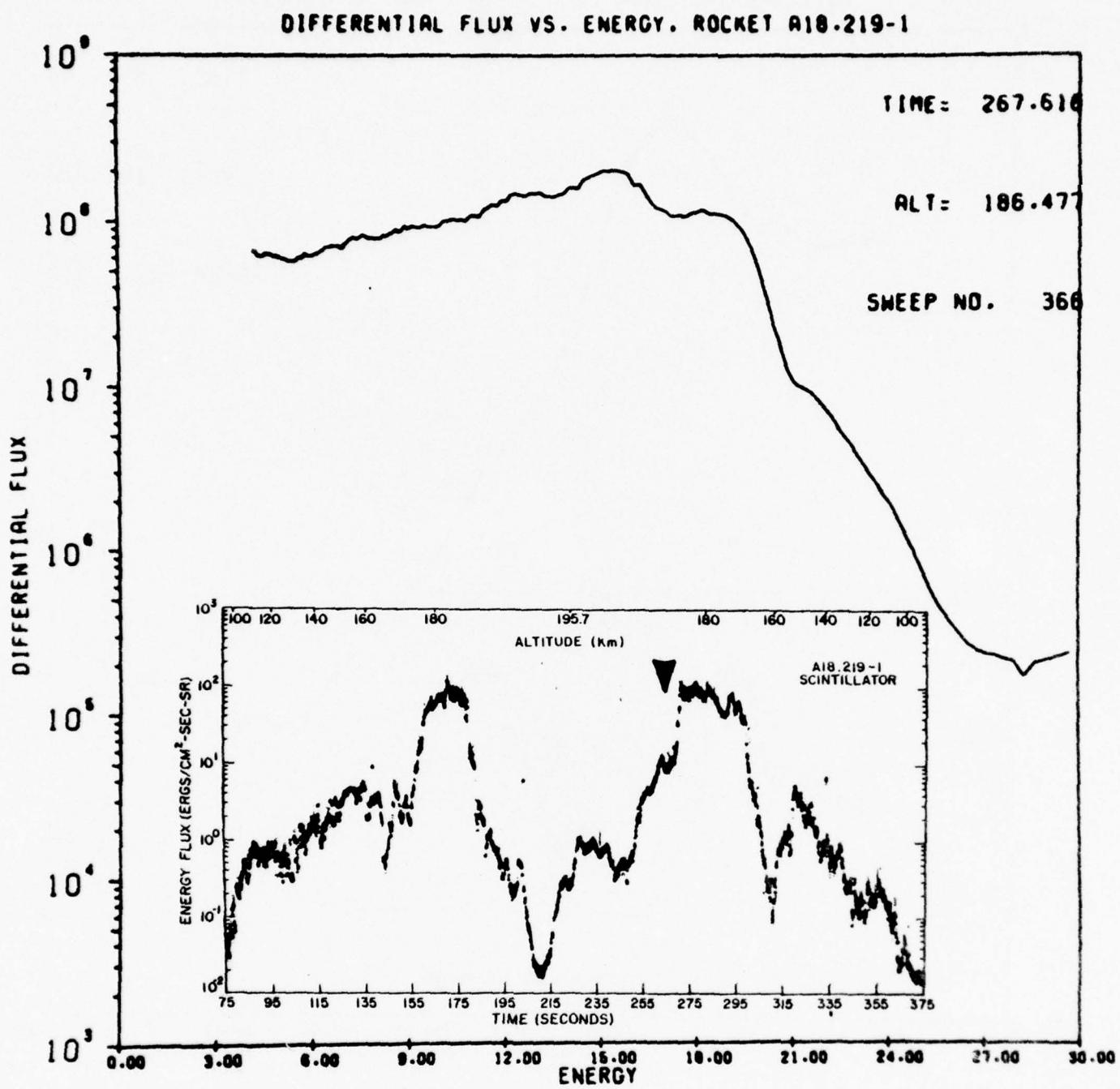


FIGURE 19

DIFFERENTIAL FLUX VS. ENERGY, ROCKET A18.219-1

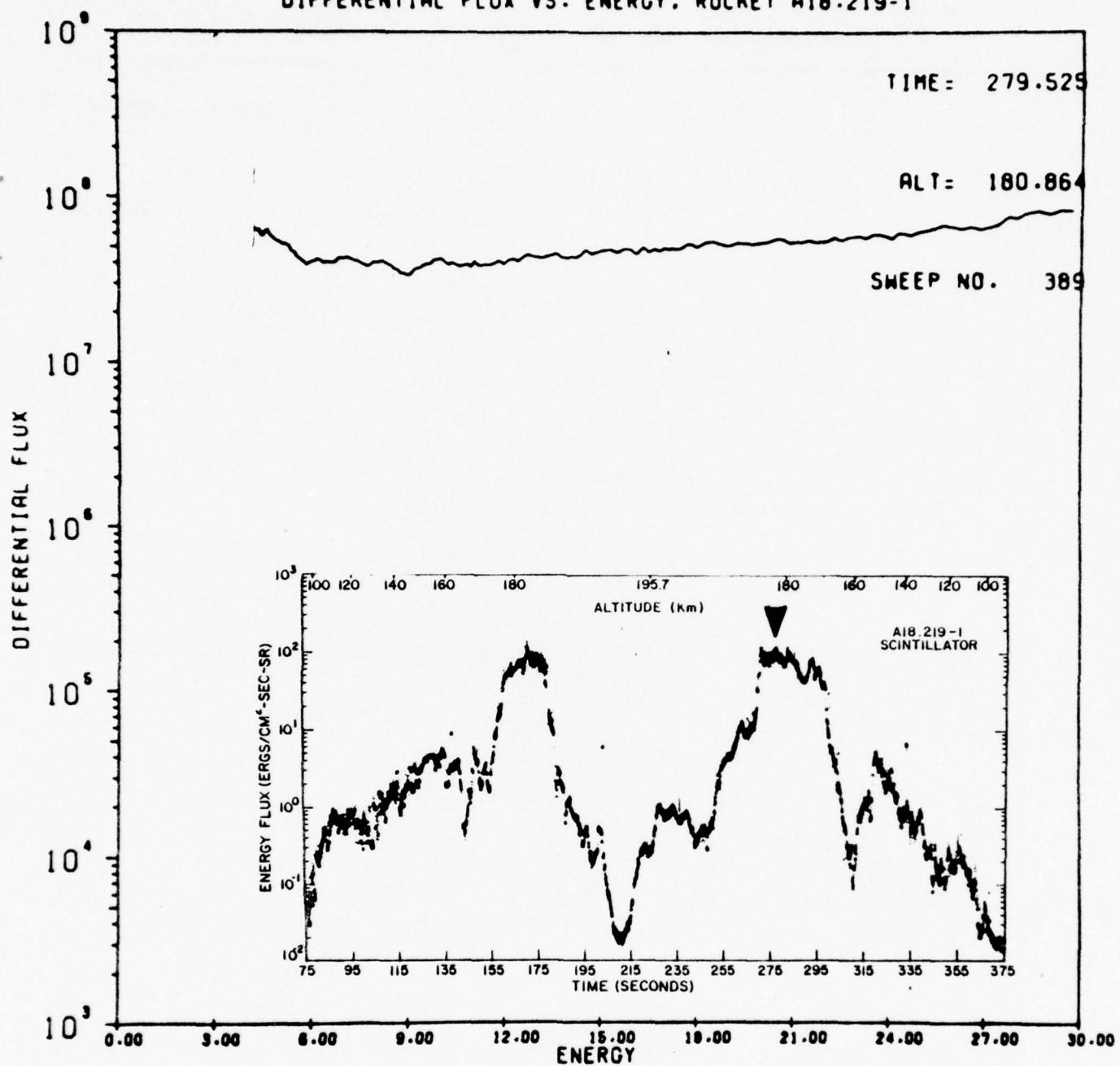


FIGURE 20

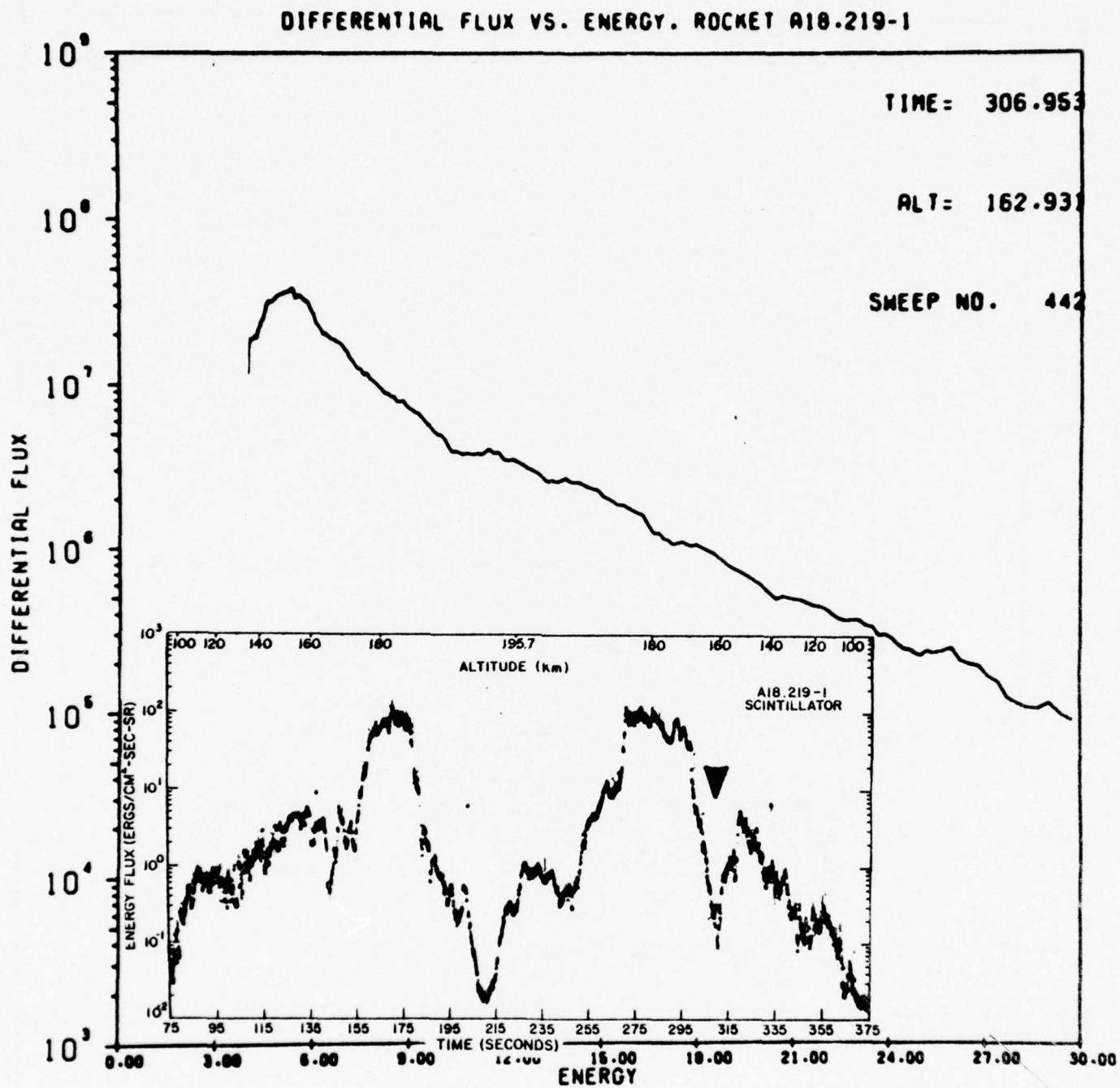


FIGURE 21

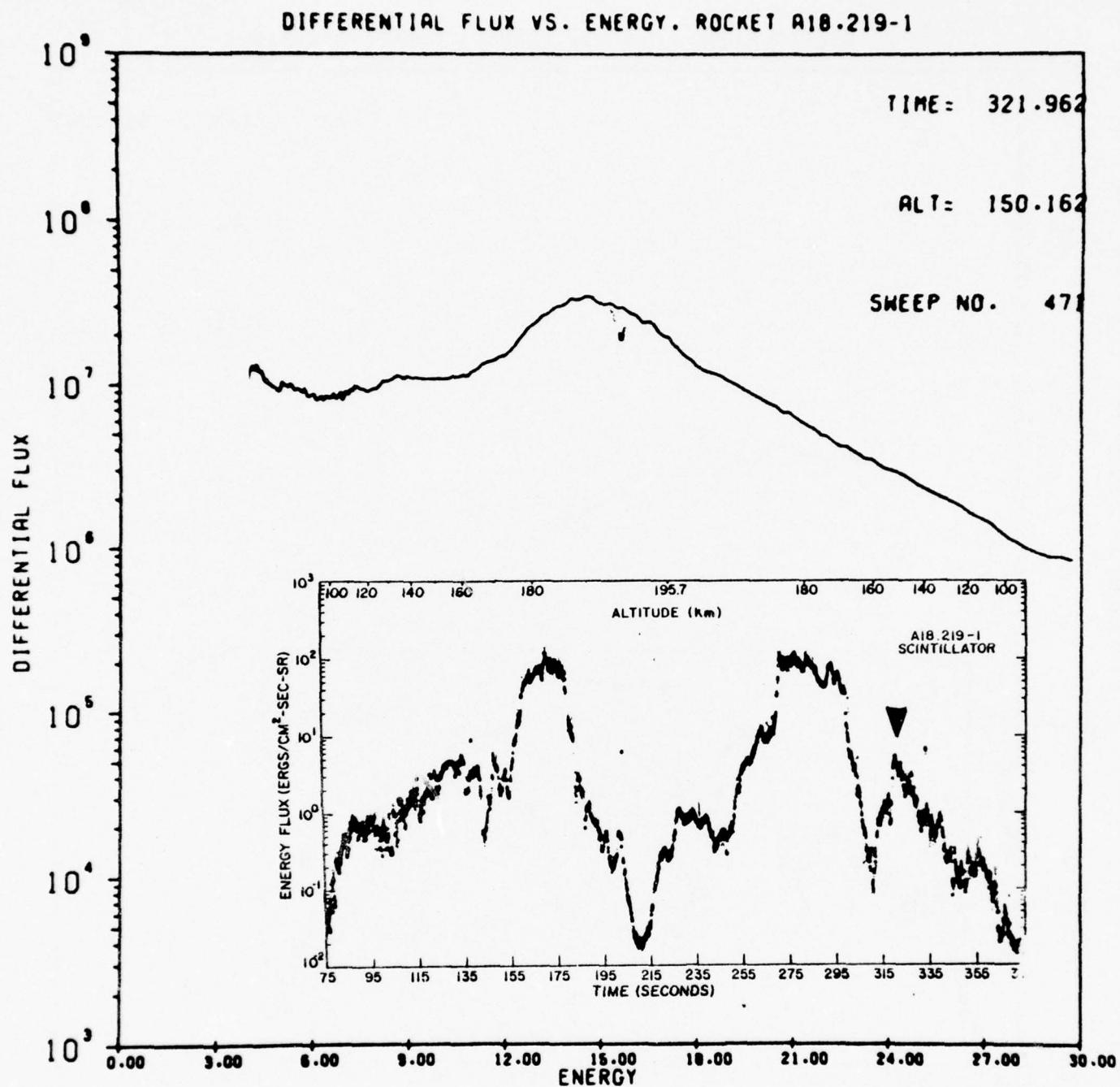


FIGURE 22

DIFFERENTIAL FLUX VS. ENERGY, ROCKET A18.219-1

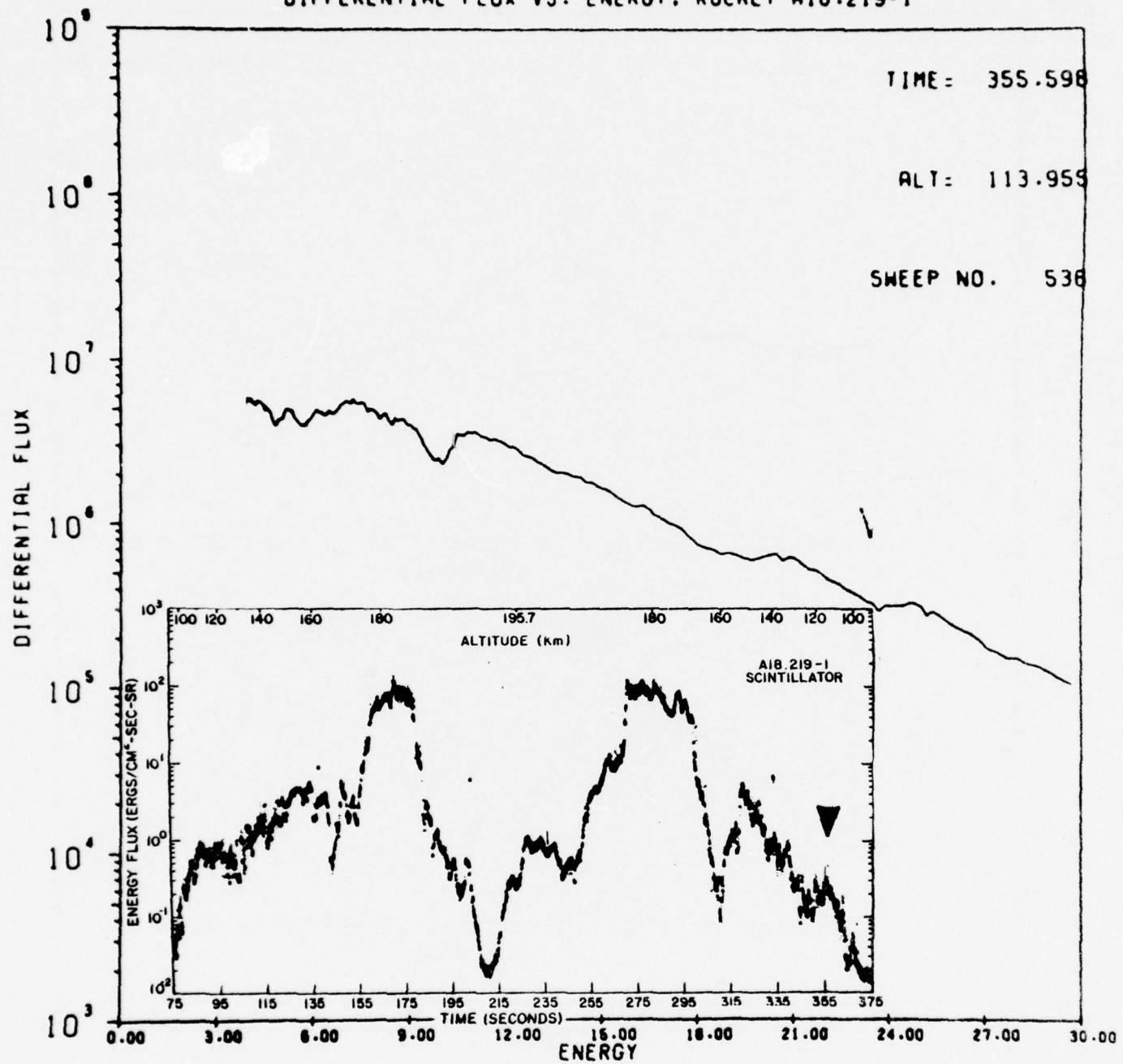


FIGURE 23

Appendix 2
Spectral Radiance Calculations at
4.3 μm , 9.6 μm , and 15 μm

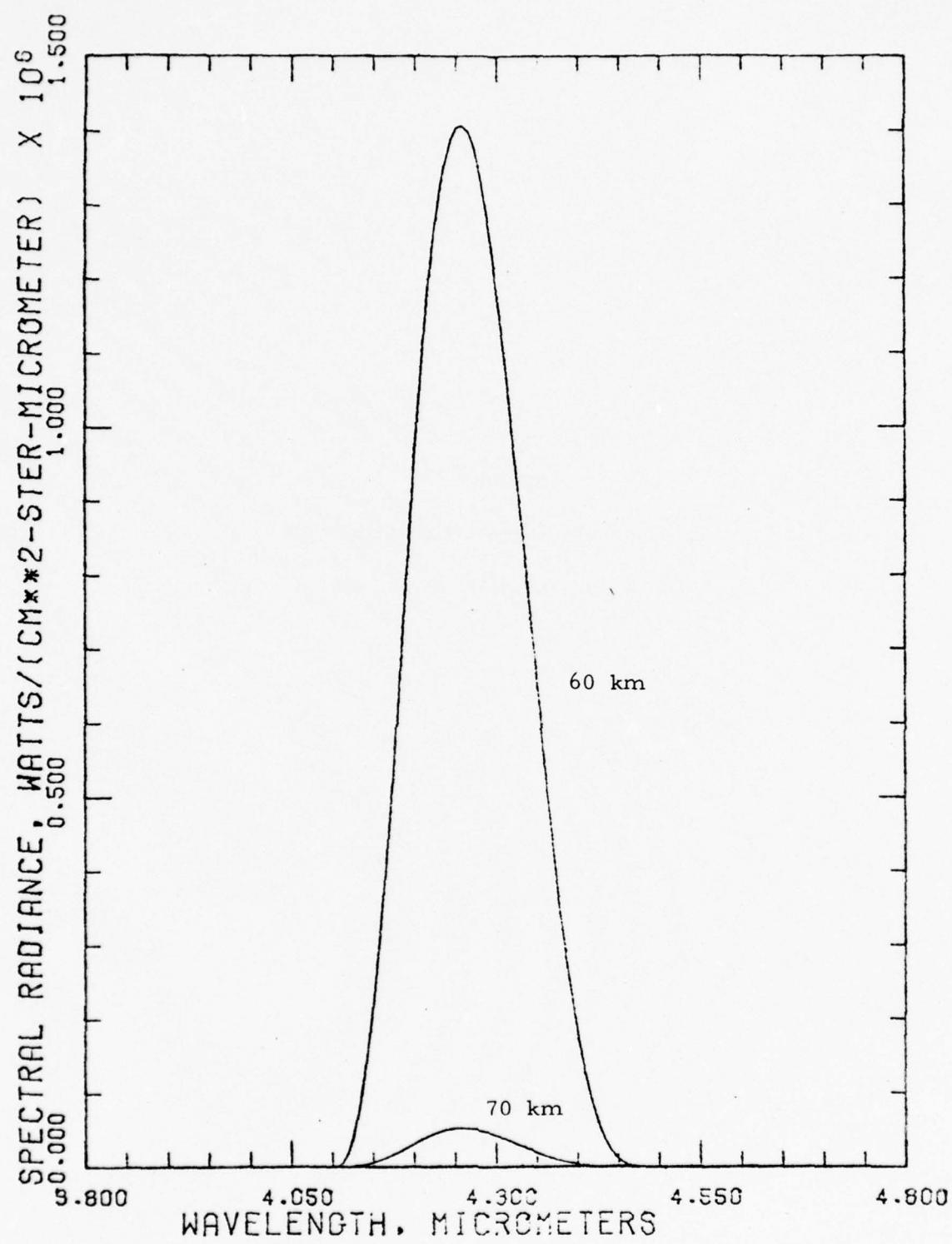


Figure A.2-1 Spectral Radiance in Upward Direction
above indicated altitude, US Standard
Atmosphere, 1962. CO₂ 4.3 μ m band.

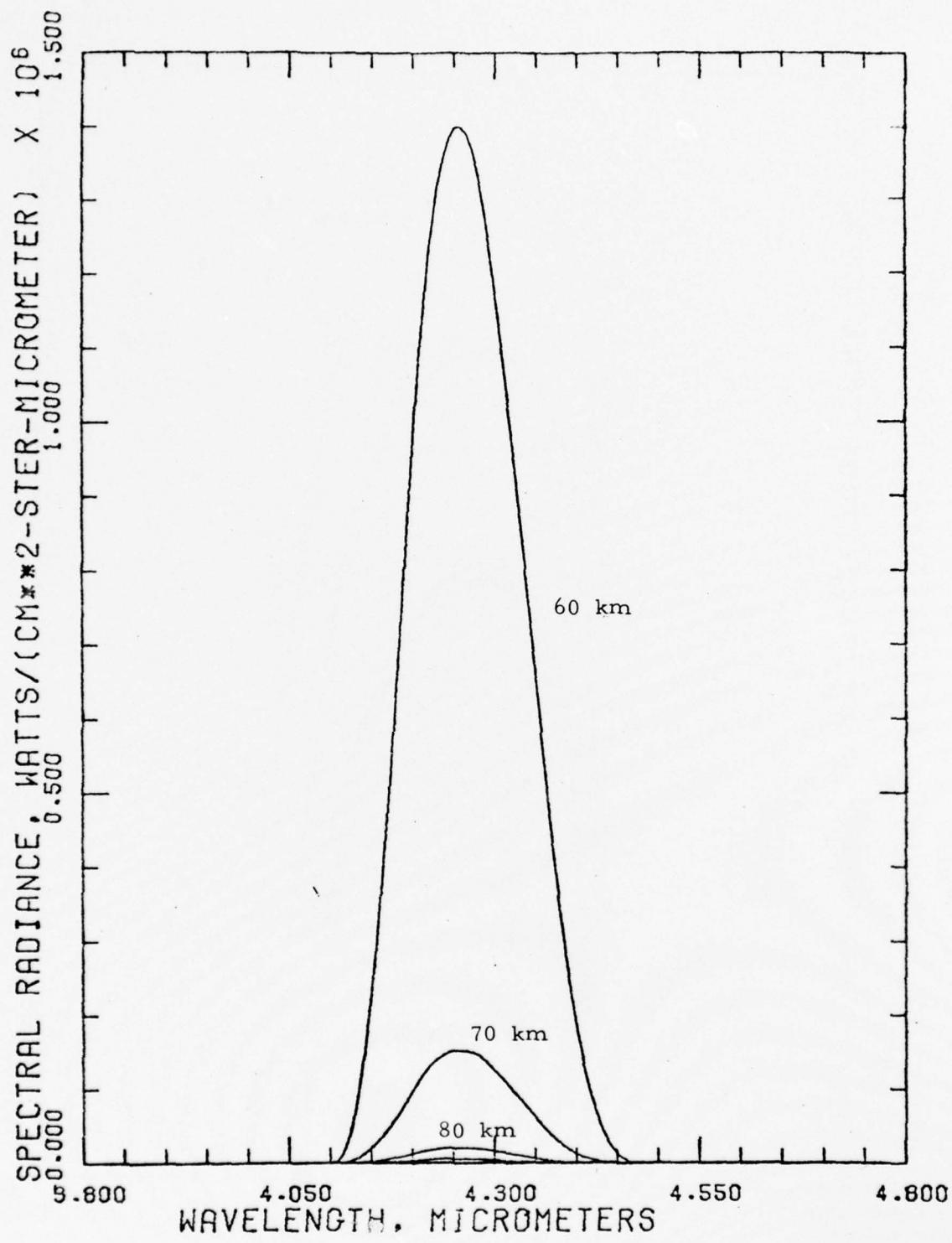


Figure A.2-2 Spectral Radiance in Upward Direction
above indicated altitude. 230 K Mesopause,
4.3 μ m CO₂ band.

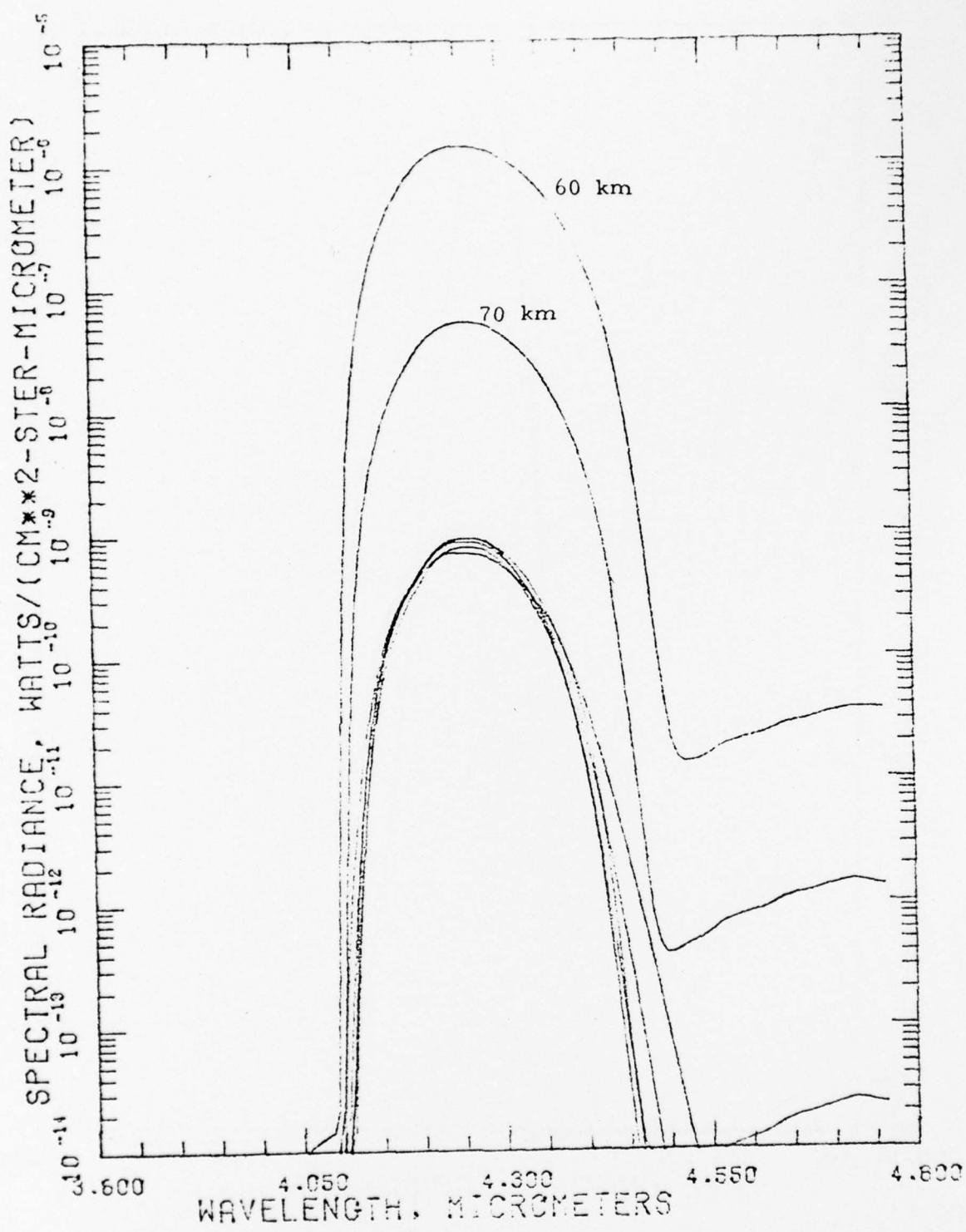


Figure A.2-3 Spectral Radiance in Upward Direction
 above indicated altitude, US Standard
 Atmosphere, 1962. CO₂ 4.3 μ m band.

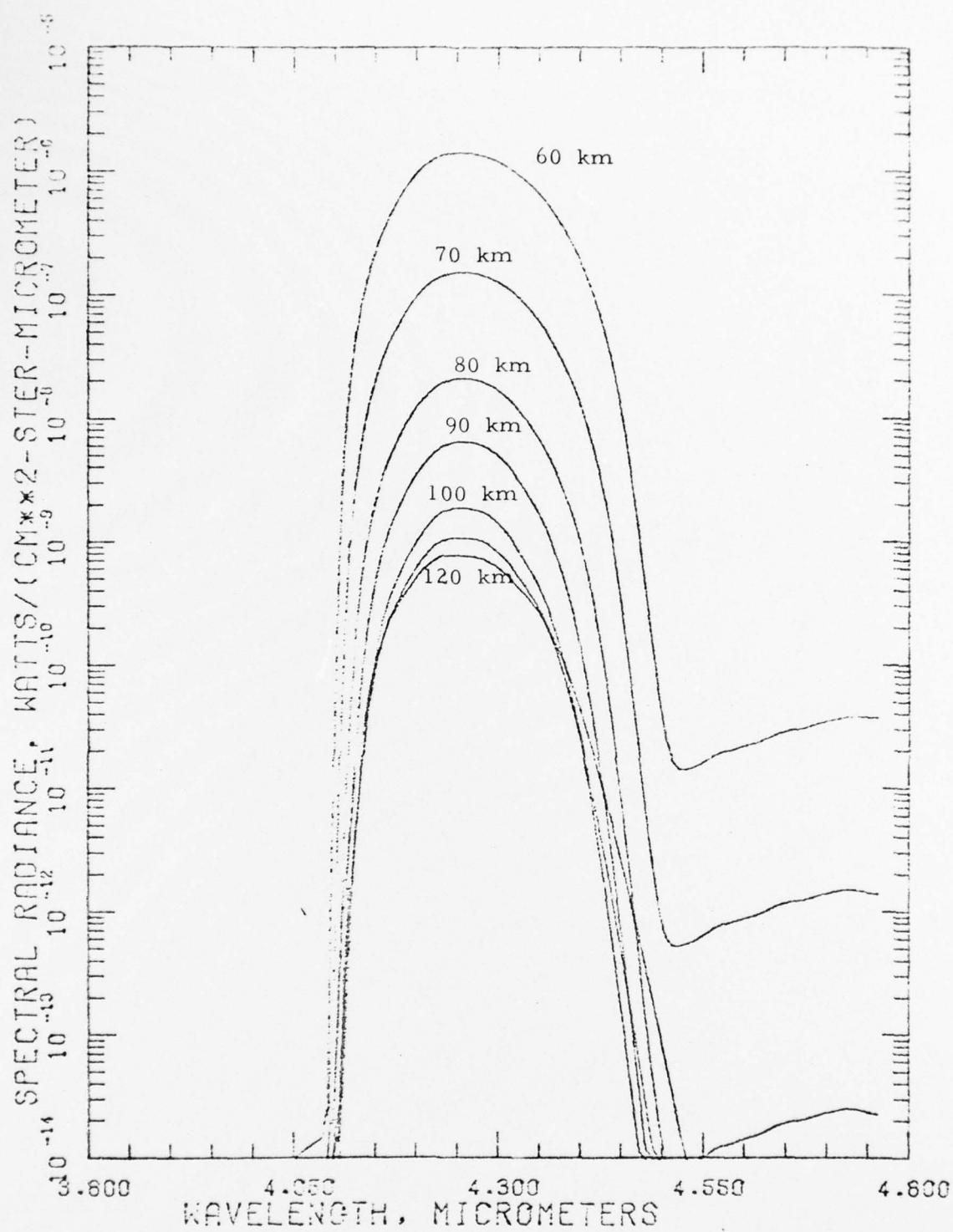


Figure A.2-4 Spectral Radiance in Upward Direction
above indicated altitude. 230 K Mesopause,
4.3 μ m CO₂ band.

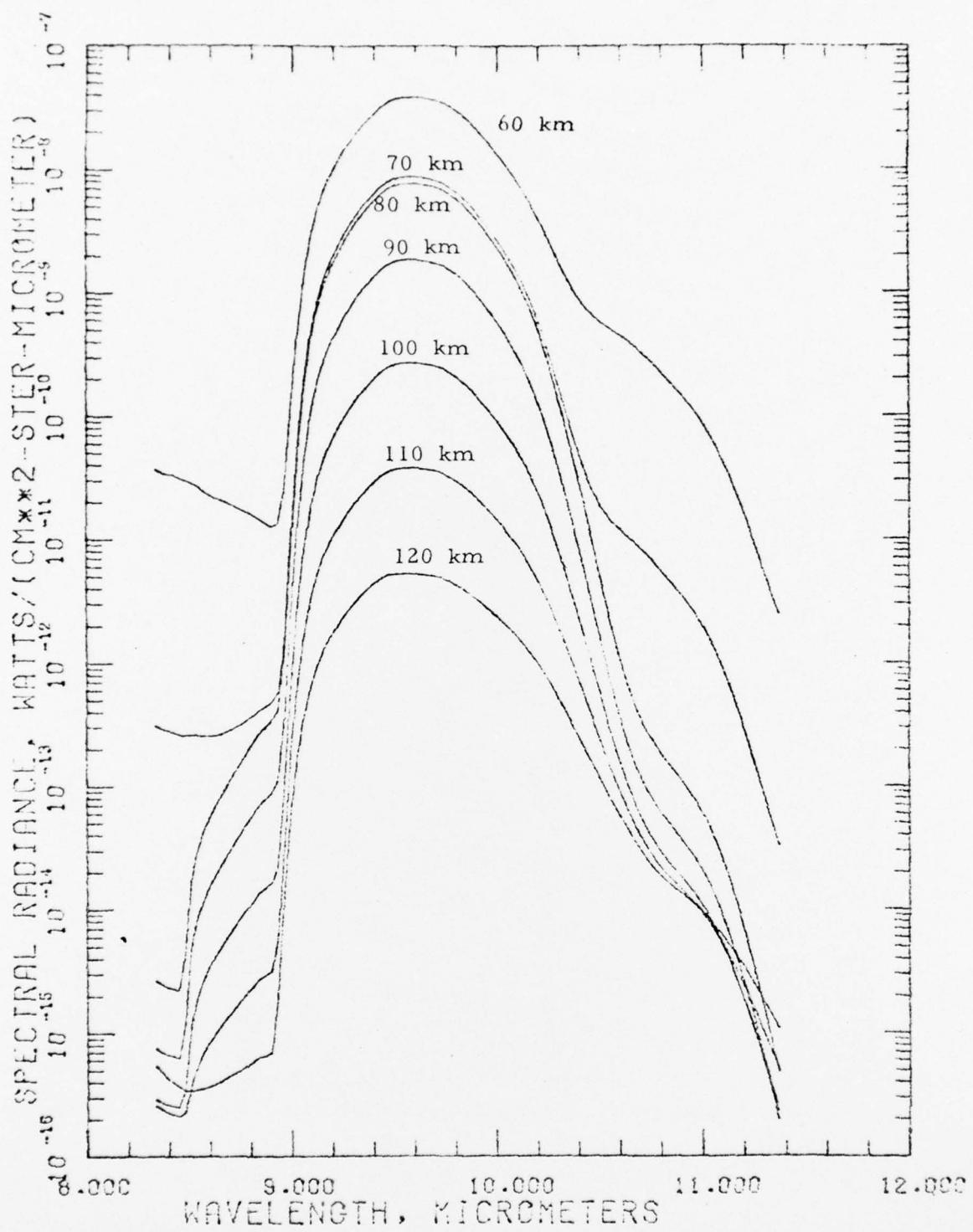


Figure A.2-5 Spectral Radiance in Upward Direction
 above indicated altitude, US Standard
 Atmosphere, 1962. Ozone 9.6 μ m band.

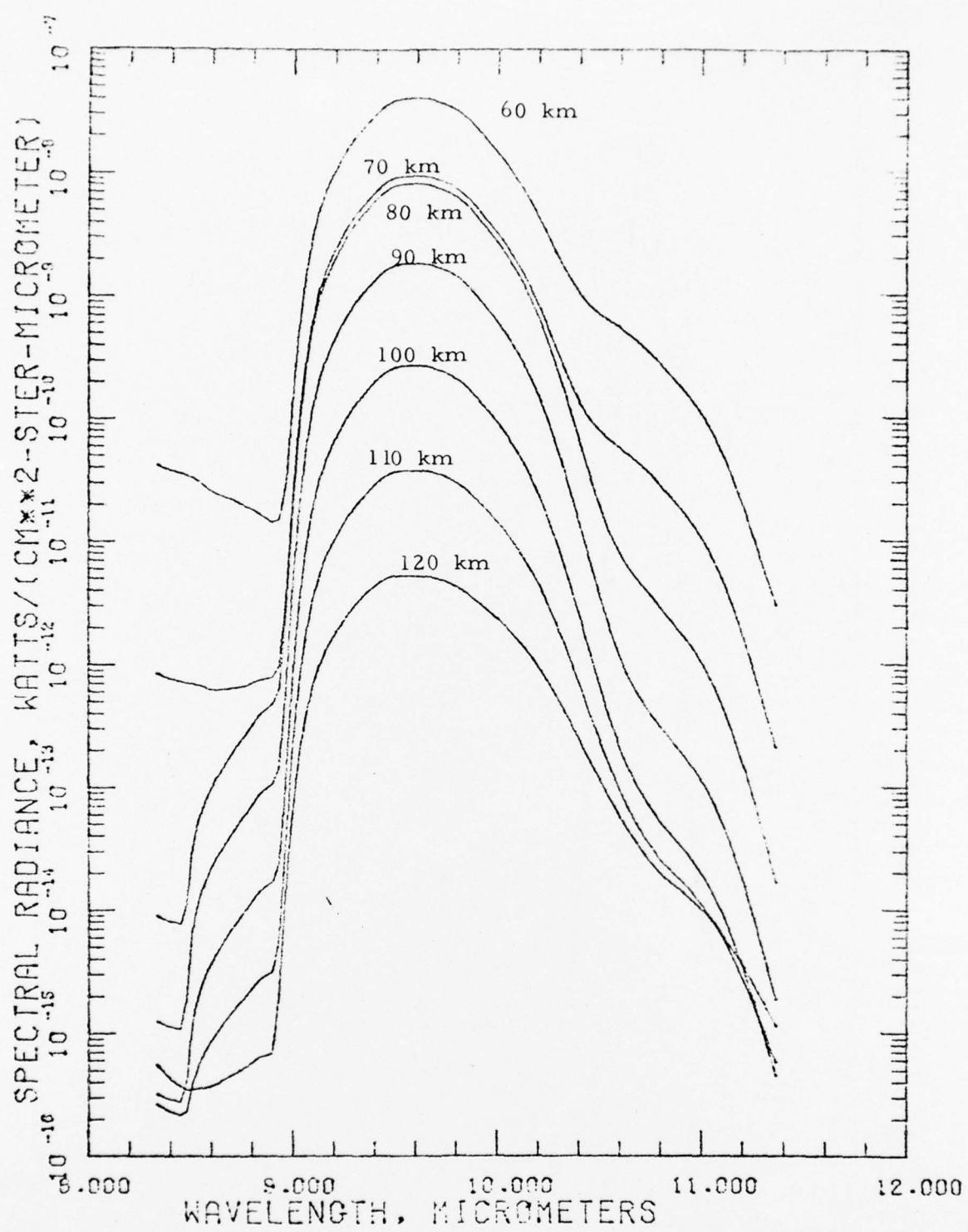


Figure A.2-6 Spectral Radiance in Upward Direction
 above indicated altitude. 230 K Mesopause,
 Ozone $9.6 \mu\text{m}$ band.

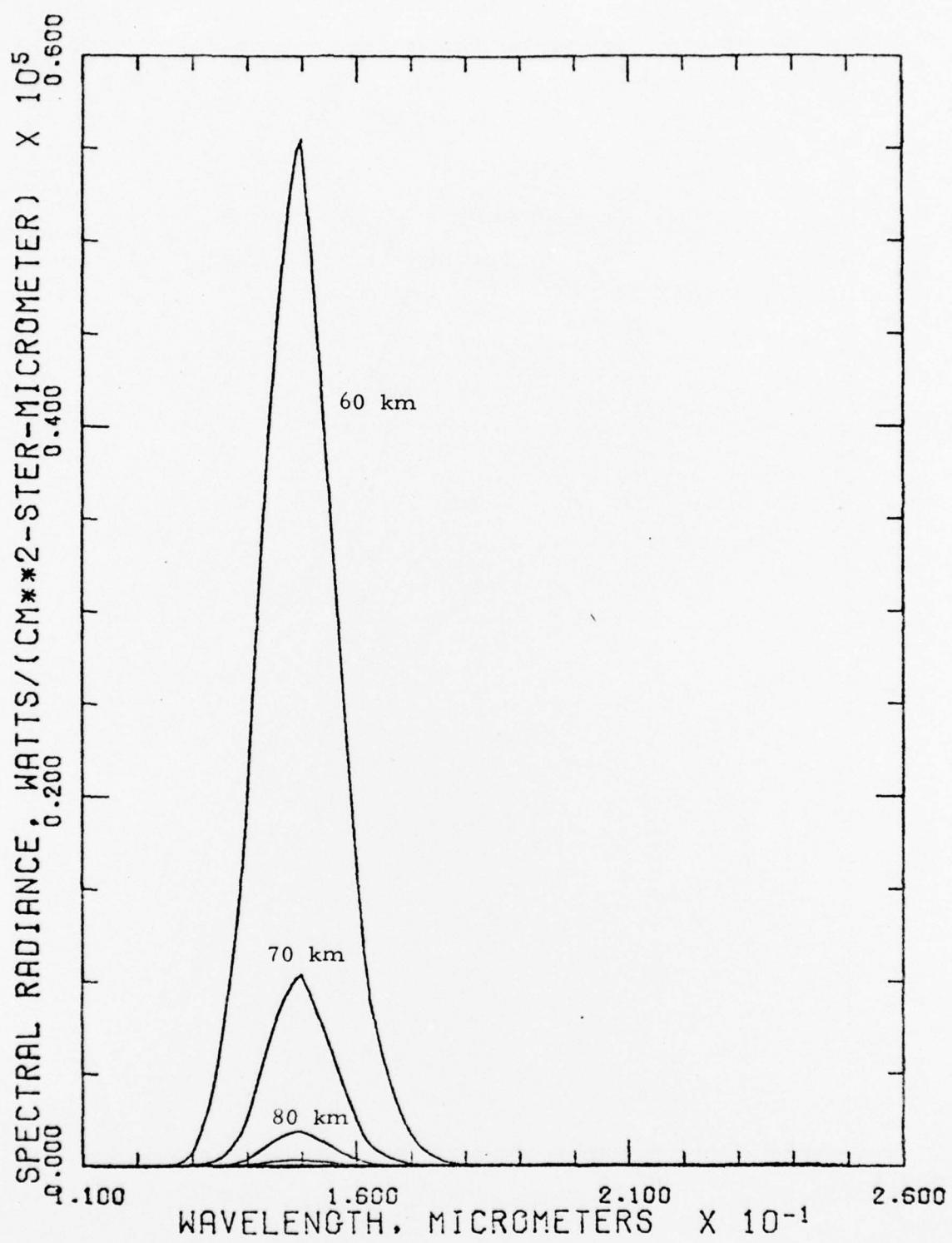


Figure A.2-7 Spectral Radiance in Upward Direction
above indicated altitude, US Standard
Atmosphere, 1962. CO_2 15 μm band.

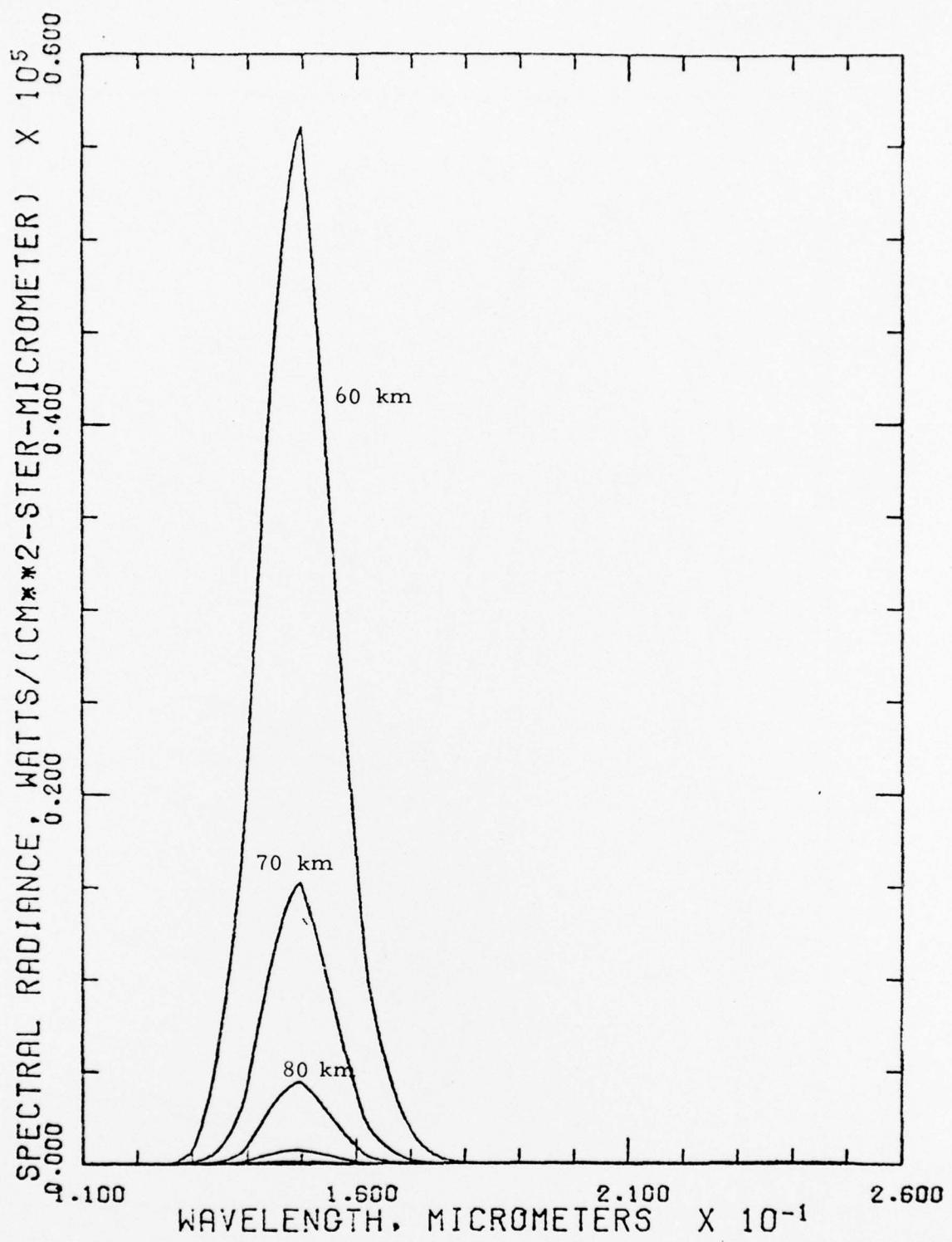


Figure A.2-8 Spectral Radiance in Upward Direction
above indicated altitude. 230 K Mesopause,
 CO_2 15μ m band.

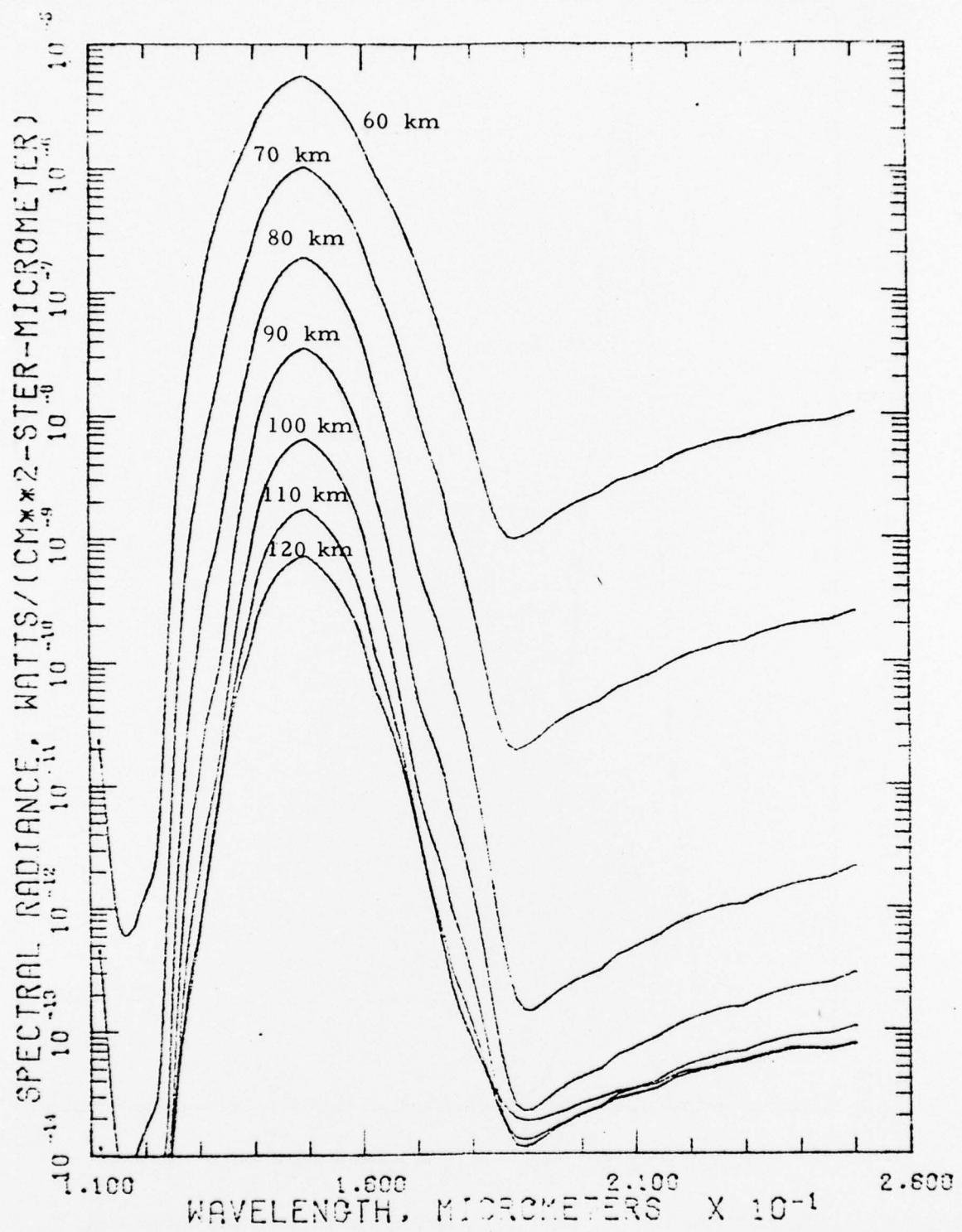


Figure A.2-9 Spectral Radiance in Upward Direction
 above indicated altitude, US Standard
 Atmosphere, 1962. CO_2 15 μm band.

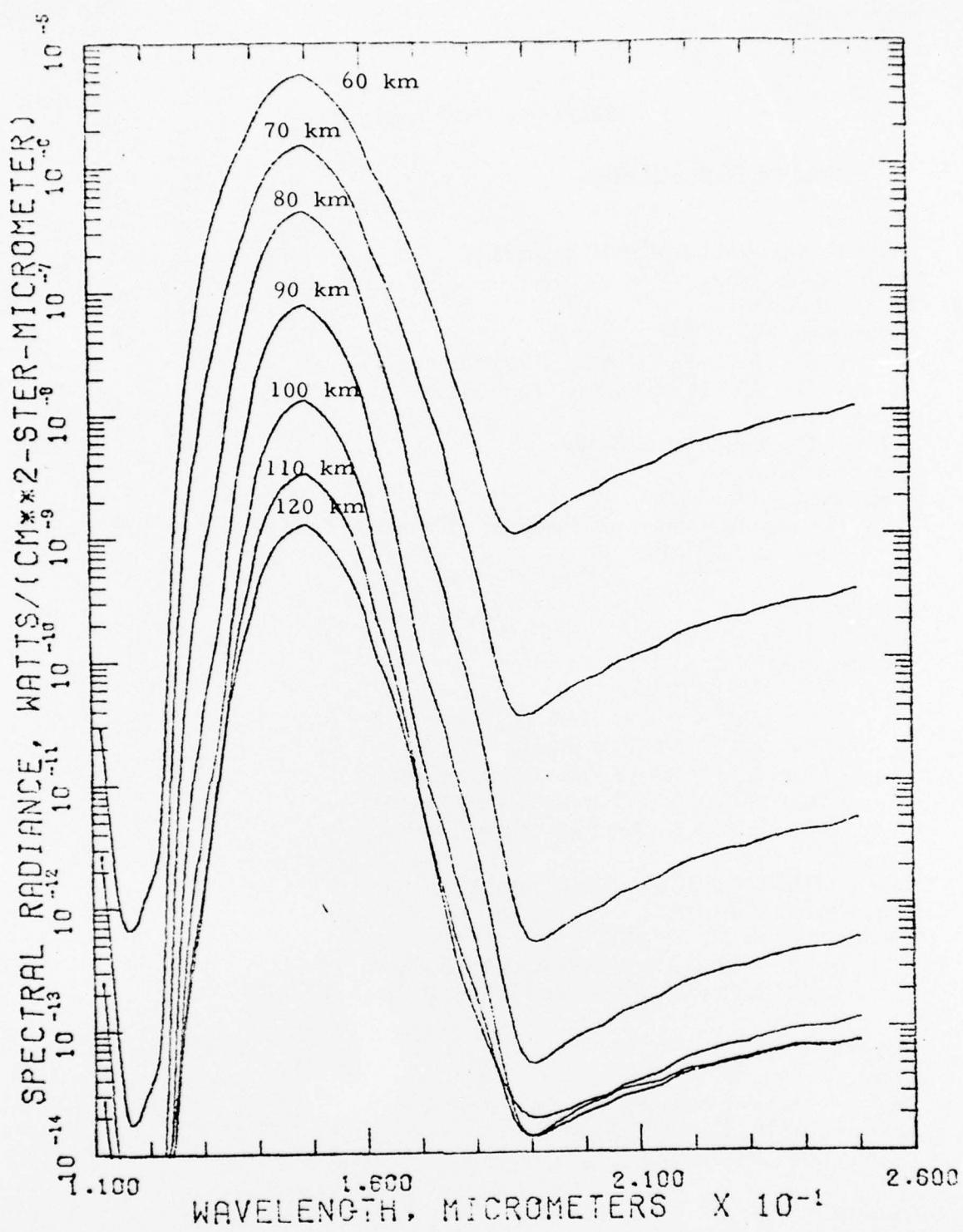


Figure A.2-10 Spectral Radiance in Upward Direction
above indicated altitude. 230 K Mesopause,
 CO_2 $15 \mu\text{m}$ band.

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